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Phase relations in the $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 join at 6 and 3.5 GPa as a model for formation of some potassium-bearing deep-seated mineral assemblages

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Abstract The join $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 has been studied at 6 GPa (890–1,500°C) and 3.5 GPa (1,000–1,100°C). K-rich melts in the join produce assemblages Cpx + Grt, Cpx + Opx, Cpx + San, and Cpx + Grt + San at 1,100–1,300°C. At $N_{\text{San}}^{\text{system}} < \sim 70$ mol%, sanidine is unstable on the solidus and appears at the liquidus, if $N_{\text{San}}^{\text{system}} > 90$ mol%. This explains a scarcity of San in mantle Cpx-rich assemblages and its association with high-K aluminosilicate melt inclusions in diamonds. In absence of San, KCpx is the only host for potassium. The K-jadeite content in KCpx systematically increases with decreasing temperature and reaches 10–12 mol% near the solidus. However, KCpx coexists with San at $N_{\text{San}}^{\text{system}} > 70$ mol% and $< 1,300^\circ\text{C}$, being formed via reaction $\text{San} + \text{L} = \text{KCpx}$. The KJd content in KCpx is controlled by the equilibrium $\text{San} = \text{KJd} + \text{SiO}_2^{\text{l}}$ that displaces to the right with increasing pressure and decreasing both the temperature and $a_{\text{SiO}_2}^{\text{l}}$. This equilibrium is considered to be responsible for the formation of San lamellae in natural UHP Cpx. In our experiments at 3.5 GPa, garnet is absent whereas the KJd and Ca-Eskola contents in Cpx are low, and the join $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 is close to binary (with the eutectic $\text{Cpx} + \text{San} + \text{L}$). Different topologies of the join at 6 and 3.5 GPa define a sequence of mineral crystallization from K-rich aluminosilicate melts during

cooling and decompression: from KCpx + Grt without San at $P > 4$ GPa to Cpx + San at $P < 4$ GPa. Similar sequence of assemblages is observed in some eclogitic xenoliths from kimberlites and Grt–Cpx rocks of the Kokchetav Complex (Northern Kazakhstan).

Abbreviations

Phases and end-members

Ab	albite ($\text{NaAlSi}_3\text{O}_8$)
Ca-Esk	Ca-Eskola pyroxene ($\text{Ca}_{0.5}\text{AlSi}_2\text{O}_6$)
Ca-Ts	Ca-Tschermack pyroxene ($\text{CaAl}_2\text{SiO}_6$)
CEn	clinoenstatite ($\text{Mg}_2\text{Si}_2\text{O}_6$)
Cos	coesite (SiO_2)
Cpx	clinopyroxene and Cpx_{ss} solid solution
Di	diopside ($\text{CaMgSi}_2\text{O}_6$)
En	orthoenstatite ($\text{Mg}_2\text{Si}_2\text{O}_6$)
Fo	forsterite (Mg_2SiO_4)
Grt	garnet and Grt_{ss} solid solution
Grs	grossular ($\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$)
Jd	jadeite ($\text{NaAlSi}_2\text{O}_6$)
KCpx	potassium-bearing clinopyroxene
KJd	potassium jadeite (KAlSi_2O_6)
Ks	kalsilite (KAlSiO_4)
Ky	kyanite (Al_2SiO_5)
L	melt
Lc	leucite (KAlSi_2O_6)
Mg-Ts	Mg-Tschermack pyroxene ($\text{MgAl}_2\text{SiO}_6$)
Opx	orthopyroxene and Opx_{ss} its solid solution
Prp	pyrope ($\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$)
Q	products of melt quenching
San	sanidine (KAlSi_3O_8)
SWd	Si-wadeite ($\text{K}_2\text{Si}_4\text{O}_9$)
Wol	orthowollastonite ($\text{Ca}_2\text{Si}_2\text{O}_6$)

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Thermodynamic symbols

P	pressure
T	temperature

N_{KJd}^{Cpx}	molar per cent of KJd in Cpx
N_{Ca}^{Grt}	molar per cent of Grs in Grt
$a_{SiO_2}^L$	silica activity in a melt
N_{San}^{system}	molar per cent of $KAlSi_3O_8$ in a system (starting mixture)

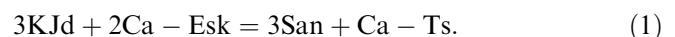
Introduction

Sanidine (San) is a very rare mineral in the mantle assemblages. Initial findings of sanidine inclusions in diamonds (e.g., Prinz et al. 1975) were attributed to the parageneses of “uncertain origin” regarding their time of the formation with respect to the host diamond. Further study showed that sanidine in diamonds occurred both as separate inclusions and polyphase inclusions with other minerals. Sanidine prefers eclogitic assemblages, e.g., omphacite, garnet, coesite, and rutile (Bulanova and Argunov 1985; Meyer and McCallum 1986; Novgorodov et al. 1990; Sobolev et al. 1998), whereas only a single its peridotitic parageneses is known (cf. Wang 1998). Sanidine is described in some eclogite xenoliths from kimberlites, i.e., in kyanite-bearing, kyanite-coesite varieties, and grosspyrites (Smyth and Hatton 1977; Wohletz and Smyth 1984; Spetsius et al. 1984; Schulze and Helmstaedt 1988; Schulze et al. 2000), as well as in biminerals (clinopyroxene and garnet, Cpx + Grt) and orthopyroxene-bearing eclogites (Mathias et al. 1970; Reid et al. 1976; Schmickler et al. 2004). The sanidine content in these rocks does not exceed 3–4 vol%. Textural relationships suggest that, in the most cases, sanidine is formed after garnet and clinopyroxene. Large grains of sanidine (Smyth and Hatton 1977; Spetsius et al. 1984) and its occurrence as inclusions in garnets (Schulze and Helmstaedt 1988; Schmickler et al. 2004) are very rare. As a rule, sanidine forms intergranular rims and small grains at the boundaries of clinopyroxenes, garnets, and other minerals. A specific appearance of sanidine in some eclogitic xenoliths are fine (less than 20 μ m) lamellae in clinopyroxenes (Reid et al. 1976; Smith et al. 1991; Schmickler et al. 2004), which are locally accompanied by lamellae of garnet, orthopyroxene, and phlogopite. Similar sanidine lamellae are also described in clinopyroxenes of the Grt–Cpx–carbonate rocks of the Kokchetav UHP Complex (Sobolev and Shatsky 1990; Perchuk et al. 1996, 2002, 2003; Zhang et al. 1997; Perchuk and Yapaskurt 1998; Shatsky and Sobolev 2003, Chap. 2.3).

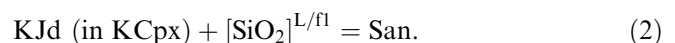
Experimental data indicate that the sanidine is stable up to pressures of 6.2–6.5 GPa in the temperature range 1,000–1,200°C (Kinomura et al. 1975; Urakawa et al. 1994; Yagi et al. 1994; Akaogi et al. 2004). At higher pressures, sanidine decomposes to the assemblage Siwadeite + kyanite + coesite (SWd + Ky + Cos) (Kinomura et al. 1975; Urakawa et al. 1994; Yagi et al.

1994; Akaogi et al. 2004). Although Ky and Cos are characteristic minerals of sanidine-bearing mantle assemblages, SWd has never been found in their parageneses. Another high-pressure precursor of sanidine might be the hollandite-structured $KAlSi_3O_8$, which is stable at pressures above 9–9.5 GPa (Ringwood et al. 1967; Kinomura et al. 1975; Urakawa et al. 1994; Yagi et al. 1994; Akaogi et al. 2004). Stachel et al. (2000b) reported a unique microcline inclusion in diamond associated with the perovskite-structured $MgSiO_3$ that could imply the initial presence of $KAlSi_3O_8$ hollandite at the time of entrapment. Davies and Harlow (2002) argue that cymrite, $KAlSi_3O_8 \cdot nH_2O$ ($n \leq 1$), might a possible precursor of sanidine at high-pressure under elevated water activity. However, cymrite is stable below 1,000°C at 5–6 GPa (e.g., Thompson et al. 1998; Davies and Harlow 2002), and therefore can be formed only at high-pressure and low-temperature metamorphic conditions (Massone et al. 2000), rather than in high-temperature mantle rocks.

Thus, sanidine could be one of the major potassium concentrators in the mantle under the anhydrous conditions, i.e., in absence of phlogopite or amphibole, at pressures below 6.5 GPa. However, it is unstable within anhydrous olivine-bearing assemblages because of reaction $San + 2Fo = Ks + 2En$ (Wendlandt and Egger 1980). Natural data indicate that sanidine is stable in eclogitic assemblages. But here, there is a much more abundant competitor of sanidine to accept potassium, namely, potassium-bearing clinopyroxene. Our recent experiments (Safonov et al. 2004) show that the clinopyroxene, crystallizing from potassic aluminosilicate and carbonate-silicate melts at pressure 5–6 GPa contains up to 2 wt% of K_2O increasing up to 5 wt% at pressure 7 GPa (Chudinovskikh et al. 2001; Bindi et al. 2002; Safonov et al. 2002, 2003). On the other hand, the presence of sanidine lamellae in deep-seated clinopyroxenes (Reid et al. 1976; Zhang et al. 1997; Perchuk et al. 1996, 2002, 2003; Schmickler et al. 2004; Shatsky and Sobolev 2003, Chap. 2.3) may suggest the decomposition of potassium-bearing clinopyroxene solid solution by decompression and cooling with the formation of sanidine via the following reaction (Luth 1997):



because KCpx may contain both the KJd and Ca-Esk end members. Perchuk and coauthors (Perchuk and Yapaskurt 1998; Perchuk et al. 2002, 2003) studied the Cpx–Grt rocks from the Kokchetav UHP complex and observed no Ca-Esk and Ca-Ts both in the lamellae-free KCpx included in garnets and K-free matrix Cpx crowded with the sanidine lamellae. On the basis of this observation, Perchuk and Yapaskurt (1998) proposed an alternative mechanism for the formation of sanidine lamellae, which is based on the reaction of KCpx with a K–Si-rich melt or fluid:



Reactions (1) and (2) imply that high-potassium clinopyroxene would be unstable within the P - T field of sanidine stability. Some experimental data (cf. Wang and Takahashi 1999) indicate indeed that the N_{KJd}^{Cpx} abruptly decreases as sanidine appears at pressures below 6.5 GPa. However, detailed experimental data on the relevant stability of sanidine and potassium-bearing clinopyroxene are absent. Because KCpx is an indicator of high potassium activity in the mantle (Harlow 1997; Perchuk et al. 2002), it is very important to know whether the assemblage of potassium-poor clinopyroxene plus sanidine replaces KCpx at the sanidine stability conditions.

The stability of sanidine in the mantle environment is directly related to the problem of origin of the mantle potassium-rich melts included in diamonds (e.g., Navon et al. 2003). A possibility of the sanidine formation from UHP potassic liquids (Wohletz and Smyth 1984; Spetsius et al. 1984; Novgorodov et al. 1990) is supported by data on occurrences of sanidine in the assemblage with inclusions of potassic aluminosilicate melts in diamonds (Bulanova and Argunov 1985; Novgorodov et al. 1990). In addition, the stoichiometry of the K-rich aluminosilicate melts that are included in diamonds is very close to $KAlSi_3O_8$ (Prinz et al. 1975; Bulanova and Argunov 1985; Novgorodov et al. 1990; Shiryayev et al. 2003; Navon et al. 2003). Experiments on the “dry” systems Di-Lc, Di-Lc-Jd, Di-San, Di-San-Jd, and Di-San-Ab at pressures above 5 GPa show that the SiO_2 -rich melts (> 50 wt%) with Al_2O_3 > 10 wt% and K_2O > 10 wt% can readily coexist with KCpx, Opx, Grs-rich Grt, Ky and Cos (Harlow 1999; Safonov et al. 2002, 2003, 2004), which are major minerals coexisting with San in the mantle assemblages.

The join Di-San is the simple system to test the idea of the formation of the San-bearing mantle assemblages from the potassium-rich melts under UHP-HP conditions and to study the relations of sanidine and potassium-bearing clinopyroxene. Detailed investigation of the join Di-San at UHP conditions has not been previously reported, although. At 6 GPa and 1,400°C, Harlow (1999) produced an assemblage of Cpx + Opx + San + K-Al-Si melt from mixtures of natural Di and San. The content of K_2O in Cpx did not exceed 0.2 wt%. Higher K_2O in Cpx (> 1 wt%) was observed by Harlow (1999) at $P > 8$ GPa in the Cpx + Grt + Ky + SWd assemblage in the sanidine-free field. Thus, the maximum KJd solubility in clinopyroxene in equilibrium with sanidine is unknown. It is still unclear whether garnet is stable in the Di-San join at $P < 8$ GPa. To expand on the results of Harlow's (1999) preliminary runs, we performed a systematic study of the join Di-San at pressures 6 and 3.5 GPa. The major purposes of our study can be formulated as follows:

1. Determination of the KJd solubility in clinopyroxene near the upper pressure boundary (e.g., 6 GPa) of the sanidine stability.
2. Testing sanidine stability in diopside-rich (e.g., eclogite-like) assemblages.
3. Examination of sanidine formation at expense of KCpx with application to sanidine lamellae in natural clinopyroxenes.
4. Tentative evaluation of a pressure boundary for Grt stability with Cpx + San + L.
5. Study of Cpx composition in equilibrium with K-rich aluminosilicate melt and/or sanidine, i.e., the KJd, Ca-Ts, Ca-Esk, and CEn contents in Cpx as a function of T and P .

Starting materials, experimental technique, and analytical procedures

Mixtures of stoichiometric gels of $CaMgSi_2O_6$ and $KAlSi_3O_8$ compositions were used as starting materials. The gels were prepared using the nitrate gelling method (Hamilton and Henderson 1968). After preparation, gels were fired to drive off volatiles and kept in desiccator. The starting gels were mixed under ethanol to produce the desired mixtures (Table 1), and about 20 mg of each mixture were placed into $Pt_{60}Rh_{40}$ capsules of 4 mm in diameter. The capsules with charges were subsequently dried during 12–24 h at 110°C and hermetically sealed by arc-welding. Despite the drying procedure, the starting gel mixtures contained small but uncontrolled amount of water, resulted in the formation of hydrous phases during quenching of melt in the low-temperature experiments (below 1,100°C).

The experiments (Table 1) were performed at 6 GPa and 890–1,500°C and 3.5 GPa and 1,000–1,100°C with the high-pressure “anvil-with-hole” apparatus (Litvin 1991; Safonov et al. 2002, 2003, 2004), which is a modification of the Bridgman-type anvil assembly (cf. Bradley 1969). The apparatus is characterized by a homogeneous distribution of pressure and temperature ($\sim 1^\circ C/mm.$) in the reaction volume of 0.1–0.15 cm³, which has been estimated by the special runs on basalt liquidus melting (see details in Litvin 1991). The high-pressure cell was prepared from limestone (Algeti mine, Georgia) individually for each run (Litvin 1991; Safonov et al. 2002, 2003). The cell is equipped with a graphite heater of 7.2 mm in length, 7.5 mm in diameter, and 0.75 mm wall thickness. The capsule was placed in the center of the cell on the holders made of pressed MgO and BN mixture (MgO : BN = 3 : 1). Pressurization of the cell was accomplished using the uni-axial compression of upper and lower anvils in a 500-ton hydraulic press. Pressure at room temperature was calibrated using bismuth phase transitions at 2.55 GPa (Bi I–Bi II), 2.7 GPa, and 7.7 GPa (Bi III–Bi V) (Homan 1975). Pressure was corrected for high temperature with the diamond-graphite curve (Kennedy and Kennedy 1976) using growth/dissolution of diamond seeds in carbon-oversaturated multicomponent carbonate melts (Spivak

Table 1 Run conditions and products of experiments in the join $\text{CaMgSi}_2\text{O}_6\text{--KAlSi}_3\text{O}_8$ at 6 and 3.5 GPa

<i>P</i> , GPa	Run number	$N_{\text{San}}^{\text{system}}$	Temperature, °C	Time, min	Run products
6	1160	10	1,350	120	Cpx, Q
	1195	10	1,150	270	Cpx, Grt, Q
	1169	20	1,450	30	Q ^b
	1168	20	1,350	120	Cpx, Q ^b
	1159	20	1,200	180	Cpx, Opx ^c , Q
	1197	20	1,150	270	Cpx, Grt ^a , Q
	1155	20	1,070	240	Cpx, Grt ^a , Q
	929	30	1,260	240	Cpx, Opx, Q
	1080	30	1,140	300	Cpx, Grt, Q
	1035	30	1,040	360	Cpx, Grt, Q
	1148	40	1,300	120	Cpx, Grt, Q
	928	40	1,260	240	Cpx, Grt, glass
	1081	40	1,140	300	Cpx, Grt, Q
	1082	40	1,040	420	Cpx, Grt, Q
	950	50	1,400	155	Cpx, glass
	926	50	1,310	240	Cpx, glass
	1006	50	1,110	360	Cpx ^c , Grt, glass
	949	50	1,040	365	Cpx, Grt, Q
	1037	50	980	840	Cpx, Grt, Q
	1036	60	1,500	20	Glass
	951	60	1,350	195	Cpx, glass
	941	60	1,160	300	Cpx, Grt, glass
	996	60	1,040	360	Cpx, Grt, glass
	927	70	1,280	240	Cpx, Grt, Q, glass
	1047	70	1,240	255	Cpx, Grt, glass
	961	70	1,130	315	Cpx, Grt, glass
	1005	70	1,110	360	Cpx, Grt, glass
	1083	70	1,050	420	Cpx, Grt, Q
	1011	80	1,380	150	Glass
	1120	80	1,340	30	Glass
	962	80	1,290	195	Cpx, Grt, glass
	946	80	1,250	240	Cpx, Grt, glass
	1121	80	1,150	240	Cpx, Grt, San, glass
	1024	80	1,090	315	Cpx, Grt, San ^c , Q, glass
	1136	80	1,000	480	Cpx, Grt, San, Q
	1093	85	1,420	60	Glass
	1109	85	1,350	30	Glass
	1086	85	1,300	120	Cpx, San, Q
	1087	85	1,200	300	Cpx, San, Q
	1090	85	1,100	360	Cpx, Grt, San, glass
	1084	90	1,320	150	San, glass
	1046	90	1,290	240	Cpx, San, Q, glass
1088	90	1,200	300	Cpx, San, Q	
1194	90	1,160	255	Cpx, San, Q	
1089	90	1,100	360	Cpx, Grt, San, glass	
978	90	890	785	Cpx, Grt, San, Q	
1149	95	1,350	60	San, glass	
1022	95	940	840	Cpx, San, SWd, Q	
3.5	1078	50	1,100	300	Cpx, Q, glass
	1071	60	1,110	240	Cpx, glass
	1073	60	1,050	360	Cpx, San, glass
	1067	70	1,100	240	Cpx, San, glass, Q
	1072	70	1,000	840	Cpx, San, Q
	1076	90	1,060	360	San, glass
	1077	95	1,100	300	San, glass [Cpx]

^aPhase is not surely identified

^bQuenching products are presented by numerous skeleton crystals of *Cpx* (Fig. 1a)

^cPhase is characterized by x-ray

and Livin 2004). As a result, run pressure was controlled to ± 0.2 GPa. Because of small reaction volume, experiments with the “anvil-with-hole” assembly cannot be monitored directly with thermocouples. Therefore, the power setting control of temperature has been used (Litvin 1991). The temperature-current power dependence of the cell was calibrated using a $\text{Pt}_{70}\text{Rh}_{30}/\text{Pt}_{94}\text{Rh}_6$ thermocouple without pressure correction. The cells were recalibrated after several runs to check the repro-

ducibility of the temperature–current power dependence. Run temperature was controlled within $\pm 20^\circ\text{C}$ using a MINITHERM-300.31 controller.

The experiments started with pressurization at ambient temperature to the desired values during 15–20 min, depending on the target pressure. Subsequently, the charge was heated to the desired temperature in a time period 10–15 min. Duration of each experiment varied in dependence on temperature (Table 1).

Table 2 Mean composition of phases produced in the join CaMgSi₂O₆-KAlSi₃O₈ at 6 and 3.5 GPa

Experiment (Table 1)	1160			1195		1159		
Phase	Cpx	glass ^(a)	Cpx ^(b)	Cpx	Grt	Cpx	Opx	
Number of points	7	6	1	10	4	25	6	
SiO ₂	55.37	79.57	55.92	54.97	41.71	55.69	59.13	
Al ₂ O ₃	0.34	8.30	1.33	1.16	22.57	1.27	1.23	
MgO	20.19	0.21	20.44	18.28	11.84	20.44	37.76	
CaO	23.96	0.97	22.26	25.36	23.97	22.43	1.73	
K ₂ O	0.04	11.00	0.00	0.11	0.01	0.28	0.06	
Total	99.90	100.05	99.95	99.88	100.10	100.11	99.91	
Formula ratio per given O								
O	6	24	6	6	12	6		
Si	1.986	10.482	1.991	1.979	3.010	1.988	1.988	
Al	0.014	1.290	0.056	0.049	1.917	0.053	0.048	
Mg	1.081	0.031	1.084	0.983	1.264	1.086	1.887	
Ca	0.920	0.136	0.849	0.981	1.860	0.855	0.062	
K	0.000	1.850	0.000	0.002	0.000	0.010	0.000	
Total	4.001	13.789	3.980	3.994	8.051	3.992	3.985	
Compositional parameters of pyroxenes and garnets ^c								
Al ^{IV}	0.014		0.009	0.021		0.012	0.002	
Al ^{M2}	0.000		0.047	0.029		0.038	0.046	
Mg ^{M2}	0.077		0.132	0.016		0.125		
□ ^{M2}	0.000		0.019	0.001		0.007		
N _{Ca} ^{Grt}					59.5			
End-members in the clinopyroxene solid solution in addition to diopside (mol%) ^d								
KJd	0.0		0.0	0.2		1.0		
Ca-Ts	1.4		0.9	2.1		1.2		
Ca-Esk	0.0		3.8	0.2		1.4		
CEn	7.7		13.2	1.6		12.5		
Experiment (Table 1)	929			1035				
Phase	Cpx	glass	Opx	Cpx	Cpx ^e	Grt	Q	
Number of points	8	1	6	20	1	10	5	
SiO ₂	55.89	71.33	57.38	55.48	51.00	44.87	55.81	
Al ₂ O ₃	2.31	13.95	1.71	1.89	10.28	23.10	19.28	
MgO	18.49	0.73	36.24	17.99	20.42	23.23	9.75	
CaO	22.09	0.57	1.70	23.66	17.87	8.75	0.09	
K ₂ O	1.23	13.36	0.17	0.92	0.40	0.00	12.25	
Total	100.01	99.94	97.20	99.94	99.97	99.95	97.18	
Formula ratio per given O								
O	6	24	6	6	6	12	24	
Si	1.998	9.638	1.982	1.992	1.800	3.084	7.963	
Al	0.097	2.220	0.070	0.080	0.428	1.871	3.241	
Mg	0.984	0.160	1.865	0.962	1.074	2.379	2.072	
Ca	0.846	0.082	0.063	0.910	0.675	0.644	0.014	
K	0.055	2.302	0.008	0.042	0.019	0.004	2.229	
Total	3.981	14.403	3.987	3.987	3.996	7.982	15.519	
Compositional parameters of pyroxenes and garnets ^c								
Al ^{IV}	0.002		0.018	0.008	0.200			
Al ^{M2}	0.095		0.052	0.072	0.228			
Mg ^{M2}	0.080			0.035	0.301			
□ ^{M2}	0.019			0.013	0.004			
N _{Ca} ^{Grt}						21.3		
End-members in the clinopyroxene solid solution in addition to diopside (mol%)								
KJd	5.5			4.2	1.9			
Ca-Ts	0.2			0.8	20.0			
Ca-Esk	3.8			2.6	0.8			
CEn	8.0			3.5	30.1			
Experiment (Table 1)	950		1006		1037			
Phase	Cpx	glass	Cpx	glass	Grt	Cpx	Grt	Q
Number of points	20	2	14	4	5	4	2	1
SiO ₂	55.31	67.36	55.20	66.45	44.05	55.04	39.58	64.62

Table 2 (Contd.)

Experiment (Table 1)	950		1006			1037			
Phase	Cpx	glass	Cpx	glass	Grt	Cpx	Grt	Q	
Al ₂ O ₃	1.61	14.73	3.49	14.62	23.86	1.97	21.97	16.04	
MgO	20.17	1.89	18.37	1.69	21.00	15.97	1.26	1.26	
CaO	22.64	1.35	22.27	1.49	11.02	24.95	36.35	2.69	
K ₂ O	0.21	14.67	0.58	15.74	0.08	1.70	0.06	12.39	
Total	99.94	100.00	99.91	99.99	100.01	99.63	99.22	97.00	
Formula ratio per given O									
O	6	24	6	24	12	6	12	24	
Si	1.976	9.271	1.969	9.223	3.048	1.995	2.980	9.110	
Al	0.068	2.389	0.147	2.392	1.945	0.084	1.949	2.664	
Mg	1.073	0.387	0.976	0.350	2.165	0.862	0.142	0.264	
Ca	0.866	0.200	0.851	0.222	0.816	0.968	2.932	0.407	
K	0.010	2.574	0.027	2.786	0.007	0.078	0.001	2.227	
Total	3.993	14.821	3.968	14.972	7.982	3.987	8.004	14.672	
Compositional parameters of pyroxenes and garnets									
Al ^{IV}	0.024		0.031			0.005			
Al ^{M2}	0.044		0.115			0.079			
Mg ^{M2}	0.117		0.091			0.000			
□ ^{M2}	0.007		0.032			0.000			
N _{Ca} ^{Grt}					27.4				
End-members in the clinopyroxene solid solution in addition to diopside (mol%)									
KJd	1.0		2.7			7.8			
Ca-Ts	2.4		3.1			0.5			
Ca-Esk	1.4		6.4			0.0			
CEn	11.7		9.1			0.0			

Experiment (Table 1)	951		996			927			
Phase	Cpx	glass	Cpx	Grt	glass	Cpx	glass	Grt	Q
Number of points	3	1	7	2	3	15	6	7	1
SiO ₂	56.06	63.47	55.47	45.17	68.04	54.84	67.85	43.42	66.30
Al ₂ O ₃	1.75	15.17	1.99	24.17	14.36	3.48	14.95	23.60	15.05
MgO	21.59	2.72	17.56	25.56	0.81	17.14	1.41	20.00	1.49
CaO	20.19	2.17	23.24	4.82	0.41	23.76	1.15	12.73	1.04
K ₂ O	0.14	16.34	1.22	0.14	16.15	0.76	15.60	0.05	15.92
Total	99.73	99.87	99.48	99.86	99.77	99.98	100.96	99.80	99.80
Formula ratio per given O									
O	6	24	6	12	24	6	24	12	24
Si	1.984	8.925	1.995	3.070	9.397	1.967	9.255	3.027	9.201
Al	0.073	2.513	0.084	1.936	2.339	0.148	2.437	1.939	2.461
Mg	1.137	0.570	0.935	2.588	0.166	0.915	0.290	2.076	0.329
Ca	0.765	0.326	0.894	0.351	0.060	0.912	0.170	0.950	0.169
K	0.006	2.929	0.056	0.012	2.844	0.035	2.752	0.022	2.818
Total	3.965	15.263	3.964	7.957	14.806	3.977	14.903	8.015	14.977
Compositional parameters of pyroxenes and garnets									
Al ^{IV}	0.016		0.005			0.033			
Al ^{M2}	0.057		0.079			0.114			
Mg ^{M2}	0.201		0.019			0.030			
□ ^{M2}	0.021		0.018			0.023			
N _{Ca} ^{Grt}				12.0				31.4	
End-members in the clinopyroxene solid solution in addition to diopside (mol%)									
KJd	0.6		5.6			3.5			
Ca-Ts	1.6		0.5			3.3			
Ca-Esk	4.2		3.6			4.6			
CEn	20.1		1.9			3.0			

Experiment (Table 1)	1005		946			1109		1084	
Phase	Cpx	glass	Grt	Cpx	glass	Grt	glass	glass	
Number of points	6	2	2	9	3	5	5	4	
SiO ₂	55.30	67.44	44.10	55.69	65.03	43.85	67.06	65.75	
Al ₂ O ₃	1.77	13.49	24.31	2.81	14.44	23.55	15.74	17.18	
MgO	18.82	2.14	23.56	16.18	1.20	20.14	2.94	1.00	

Table 2 (Contd.)

Experiment (Table 1)	1005			946			1109	1084
Phase	Cpx	glass	Grt	Cpx	glass	Grt	glass	glass
CaO	23.54	2.07	8.00	23.50	1.33	12.15	3.43	1.32
K ₂ O	0.35	14.63	0.02	1.55	17.92	0.25	15.23	14.74
Total	99.78	99.77	100.00	99.73	99.91	99.94	104.40	99.99
Formula ratio per given O								
O	6	24	12	6	24	12	24	24
Si	1.984	9.319	3.026	2.000	9.164	3.049	8.956	9.060
Al	0.075	2.197	1.966	0.119	2.397	1.930	2.477	2.790
Mg	1.006	0.441	2.409	0.867	0.252	2.086	0.585	0.205
Ca	0.905	0.306	0.588	0.906	0.200	0.905	0.491	0.195
K	0.016	2.577	0.002	0.071	3.222	0.022	2.593	2.590
Total	3.986	14.840	7.992	3.963	15.236	7.992	15.102	14.840
Compositional parameters of pyroxenes and garnets								
Al ^{IV}	0.016			0.000				
Al ^{M2}	0.059			0.119				
Mg ^{M2}	0.065			0.000				
□ ^{M2}	0.014			0.023				
N _{Ca} ^{Grt}			19.6			30.3		
End-members in the clinopyroxene solid solution in addition to diopside (mol%)								
KJd	1.6			7.1				
Ca-Ts	1.6			0.0				
Ca-Esk	2.8			4.6				
CEn	6.5			0.0				

Experiment (Table 1)	1024			1086				
Phase	Cpx	glass	Grt	San	Q	Cpx	San	Q
Number of points	17	8	8	9	3	12	5	3
SiO ₂	55.65	68.14	44.12	64.39	66.78	55.34	65.30	63.09
Al ₂ O ₃	2.08	14.71	23.74	17.87	15.14	0.00	18.28	15.50
MgO	18.04	0.74	21.56	0.02	0.59	18.61	0.00	0.67
CaO	22.92	0.41	10.31	0.02	0.24	26.13	0.07	1.01
K ₂ O	1.03	15.97	0.18	17.67	17.26	0.20	16.84	15.71
Total	99.73	99.98	99.92	99.96	100.01	100.30	100.49	95.98
Formula ratio per given O								
O	6	24	12	8	24	6	8	24
Si	1.999	9.397	3.051	3.000	9.293	1.994	3.007	9.139
Al	0.088	2.391	1.934	0.981	2.482	0.000	0.992	2.647
Mg	0.965	0.152	2.221	0.001	0.122	0.999	0.000	0.144
Ca	0.881	0.061	0.764	0.001	0.036	1.008	0.003	0.156
K	0.047	2.809	0.016	1.050	3.063	0.009	0.989	2.903
Total	3.981	14.814	7.996	5.037	14.997	4.011	4.991	14.989
Compositional parameters of pyroxenes and garnets								
Al ^{IV}	0.001					0.006		
Al ^{M2}	0.087					0.000		
Mg ^{M2}	0.052					0.000		
□ ^{M2}	0.019					0.000		
N _{Ca} ^{Grt}			25.6					
End-members in the clinopyroxene solid solution in addition to diopside (mol%)								
KJd	4.7					0.0		
Ca-Ts	0.1					0.0		
Ca-Esk	3.8					0.0		
CEn	5.2					0.0		

Experiment (Table 1)	1087 ^f				1088		
Phase	Di	KCpx	San	Q	Cpx	San	Q
Number of points	19	9	7	2	9	4	4
SiO ₂	55.45	54.56	65.44	66.12	55.45	65.82	64.86
Al ₂ O ₃	0.03	4.04	18.29	15.39	0.00	18.27	14.54
MgO	18.98	16.92	0.01	0.89	18.55	0.03	0.54
CaO	25.41	23.59	0.03	1.08	26.62	0.03	1.11
K ₂ O	0.10	0.91	16.33	16.55	0.17	17.30	16.18
Total	99.98	100.02	100.10	100.01	100.79	101.44	97.23

Table 2 (Contd.)

Experiment (Table 1)	1087 ^f				1088		
Phase	Di	KCpx	San	Q	Cpx	San	Q
Formula ratio per given O							
O	6	6	8	24	6	8	24
Si	1.998	1.957	3.014	9.201	1.991	3.008	9.281
Al	0.001	0.171	0.993	2.523	0.000	0.984	2.451
Mg	1.019	0.903	0.000	0.183	0.992	0.002	0.116
Ca	0.981	0.906	0.002	0.161	1.023	0.001	0.169
K	0.006	0.042	0.959	2.937	0.008	1.008	2.953
Total	4.004	3.979	4.969	15.005	4.014	5.004	14.970
Compositional parameters of pyroxenes and garnets							
Al ^{IV}	0.002	0.043			0.009		
Al ^{M2}	0.000	0.128			0.000		
Mg ^{M2}	0.019	0.031			0.000		
□ ^{M2}	0.000	0.021			0.000		
N _{Ca} ^{Grt}							
End-members in the clinopyroxene solid solution in addition to diopside (mol%)							
KJd	0.0	4.2			0.0		
Ca-Ts	0.0	4.3			0.0		
Ca-Esk	0.0	4.2			0.0		
CEn	1.9	3.1			0.0		

Run sample (Table 1)	1090 ^f							1046			
Phase	KCpx	Al-Cpx	Di	glass	Grt	Ca-Grt	San	Q ^g	Cpx	San	glass
Number of points	2	14	18	7	11	4	5	1	5	3	5
SiO ₂	55.26	52.56	55.33	56.07	42.65	42.15	65.14	70.01	55.59	64.64	65.80
Al ₂ O ₃	2.02	4.97	0.01	17.05	23.34	22.67	18.36	11.25	3.00	17.88	14.94
MgO	17.46	16.62	18.80	1.48	17.67	12.39	0.00	0.26	17.99	0.11	0.64
CaO	24.30	25.62	25.76	2.62	16.29	22.60	0.13	1.31	21.69	0.04	0.58
K ₂ O	0.93	0.22	0.08	22.76	0.05	0.06	16.34	17.15	1.25	16.98	18.16
Total	99.97	99.99	99.98	99.98	99.98	99.86	99.97	99.98	99.52	99.65	100.12
Formula ratio per given O											
O	6	6	6	24	12	12	8	24	6	8	24
Si	1.974	1.887	2.000	8.504	3.005	3.051	3.032	9.733	1.991	3.008	9.232
Al	0.108	0.257	0.000	3.037	1.937	1.873	0.958	1.843	0.127	0.980	2.462
Mg	0.913	0.862	1.008	0.187	1.854	1.380	0.000	0.054	0.960	0.008	0.134
Ca	0.954	0.972	0.993	0.155	1.229	1.702	0.019	0.195	0.832	0.002	0.087
K	0.048	0.015	0.000	4.189	0.005	0.015	0.960	3.041	0.057	1.008	3.254
Total	3.996	3.992	4.000	16.072	8.029	8.020	4.969	14.866	3.967	5.006	15.169
Compositional parameters of pyroxenes and garnets											
Al ^{IV}	0.026	0.113	0.000						0.009		
Al ^{M2}	0.082	0.144	0.000						0.118		
Mg ^{M2}	0.000	0.013	1.000						0.079		
□ ^{M2}	0.000	0.008	0.000						0.033		
N _{Ca} ^{Grt}					39.9	55.2					
End-members in the clinopyroxene solid solution in addition to diopside (mol%)											
KJd	4.8	1.5	0.0						5.7		
Ca-Ts	2.6	11.3	0.0						0.9		
Ca-Esk	0.0	1.6	0.0						6.6		
CEn	0.0	1.3	0.0						7.9		

Experiment (Table 1)	1022			978		
Phase	Cpx ^h	San	SWd	Cpx ^h	Grt	Q
Number of points	1	3	3	1	3	3
SiO ₂	56.63	66.97	72.63	55.91	39.88	64.69
Al ₂ O ₃	4.45	18.55	5.79	3.54	21.82	14.92
MgO	14.47	0.02	0.85	15.32	4.52	0.10
CaO	20.36	0.04	0.52	22.50	33.54	0.48
K ₂ O	2.62	14.97	23.71	2.72	0.08	16.49
Total	98.53	100.56	103.50	99.99	101.84	96.68

Table 2 (Contd.)

Experiment (Table 1)	1022			978		
Phase	Cpx ^h	San	SWd	Cpx ^h	Grt	Q
O	6	8	9	6	12	24
Si	2.028	3.038	3.791	2.012	2.972	9.301
Al	0.188	0.991	0.356	0.150	1.916	2.528
Mg	0.772	0.002	0.066	0.821	0.502	0.021
Ca	0.781	0.002	0.029	0.867	2.677	0.074
K	0.119	0.866	1.578	0.125	0.008	3.024
Total	3.887	4.899	5.820	3.975	8.075	14.948
Compositional parameters of pyroxenes and garnets						
Al ^{IV}	0.000			0.000		
Al ^{M2}	0.188			0.150		
Mg ^{M2}	0.000			0.000		
□ ^{M2}	0.100			0.008		
N _{Ca} ^{Grt}					84.2	
End-members in the clinopyroxene solid solution in addition to diopside (mol%)						
KJd	11.9			12.5		
Ca-Ts	0.0			0.0		
Ca-Esk	20.0			1.6		
CEn	0.0			0.0		
Experiment (Table 1)						
	1067			1072		
Phase	glass	Cpx	San	Q	Cpx ^h	San
Number of points	5	9	3	2	3	7
SiO ₂	65.74	55.35	65.07	72.47	57.28	65.18
Al ₂ O ₃	16.00	1.10	18.15	10.70	1.26	17.91
MgO	1.27	20.18	0.00	0.50	18.29	0.31
CaO	1.04	23.21	0.00	0.10	22.58	0.45
K ₂ O	15.80	0.17	16.76	12.43	0.61	16.12
Total	99.85	100.01	99.98	96.2	100.01	99.97
Formula ratio per given O						
O	24	6	8	24	6	8
Si	9.128	1.981	3.010	10.075	2.038	3.009
Al	2.618	0.046	0.989	1.753	0.053	0.974
Mg	0.263	1.076	0.000	0.104	0.970	0.021
Ca	0.155	0.889	0.000	0.015	0.861	0.024
K	2.798	0.007	0.990	2.204	0.028	0.949
Total	14.962	3.999	4.989	14.151	3.949	4.978
Compositional parameters of pyroxenes						
Al ^{IV}		0.019			0.000	
Al ^{M2}		0.027			0.053	
Mg ^{M2}		0.103			0.039	
□ ^{M2}		0.000			0.073	
End-members in the clinopyroxene solid solution in addition to diopside (mol%)						
KJd		0.7			2.8	
Ca-Ts		1.9			0.0	
Ca-Esk		0.0			14.6	
CEn		10.3			3.9	

^aExperiment 1160—residual glass in quench

^bExperiment 1160—dendritic clinopyroxene in quench

^cComponents in the sites of clinopyroxenes were calculated using the following scheme: Al^{IV} = 2 - (Si), if (Si) ≤ 2 a.p.f.u., and Al^{IV} = 0, if (Si) > 2 a.p.f.u. (i.e., assuming that Si occupies exclusively the tetrahedral site, while the Si content above 2.0 a.p.f.u. results from poor-quality analyses of tiny crystals); Al^{VI} = (Al) - Al^{IV}; Mg^{M1} = 1 - Al^{M1}; Mg^{M2} = (Mg) - Mg^{M1}; □^{M2} = 1 - (Ca) - (K) - Mg^{M2}, where elements in parentheses denote atoms in formula unit based on 6 oxygen atoms. Molar per cent of grossular component in garnet is calculated with following formula: N_{Ca}^{Grt} = Ca/(Ca + Mg), where elements denote atoms in formula unit based on 12 oxygen atoms

^dEnd-members (mol%) of clinopyroxene solid solution were calculated using the following procedure: KJd = (K)*100, if (K) > (Al) - Al^{IV}, unless KJd is assumed to be zero; Ca-Ts = Al^{IV}*100; CEn = Mg^{M2}*100; Ca-Esk = (2*□^{M2})*100

^eInclusion in garnet

^fCompositional types of clinopyroxenes and garnets in the experiments 1087 and 1090

^gQuench in the sanidine-clinopyroxene cumulates (see lower inset in Fig. 1e)

^hThe Si content above 2.0 a.p.f.u. and high Ca-Esk contents result from poor-quality analyses of acicular crystals

Abbreviations: KCpx, potassium-bearing clinopyroxene; Al-Cpx, Ca-Ts-rich clinopyroxene at contacts with garnet; Di, KJd-free clinopyroxene; Grt, centers of garnet grains; Ca-Grt, garnet at the contact with clinopyroxene

Experiments were quenched by shutting power off and subsequently depressurized.

Each run sample was divided into several pieces. This procedure helps to examine thoroughly textural features of run products, which are texturally heterogeneous in some cases (see below). The pieces were embedded in epoxy and polished. After preliminary examination in reflected light, the microscopic features of run products and phase compositions (Table 2) were studied by means of BSE with a CamScan MV2300 (VEGA TS 5130MM) electron microscope equipped with the EDS electron microprobe Link INCA Energy at the RSMA laboratory of the Institute of Experimental Mineralogy. Some samples were examined twice or thrice after repeated polishing. Microprobe analyses were performed at 20 kV accelerating potential, 10 nA of beam current and beam diameter of 3 μm . Microprobe analyses of glass (or the products of quenching) were made using a defocused beam or scanning of area $10 \times 10 \mu\text{m}$. The following standards were used: synthetic SiO_2 for Si, synthetic MgO for Mg, synthetic Al_2O_3 for Al, natural wollastonite for Ca, and natural microcline for K. The ZAF matrix correction was used. No normalization procedure is applied. Mineral formulae (Table 2) were calculated on the basis of a specified number of oxygen atoms.

Hand-picked crystal fragments of Opx from the run product 1159, and Cpx from the run products 1006 and 927, were used for the X-ray single-crystal diffraction studies. Intensity data were collected on an automated diffractometer (Enraf Nonius—CAD4) using Mo $K\alpha$ radiation monochromatized by a flat graphite crystal. Both the intensity corrections and the strategies of the crystal structure refinements were similar to those carried out in our previous studies (Bindi et al. 2002, 2003). Unit-cell parameters of Opx (space group *Pbca*), determined by centering 25 high- θ ($20\text{--}27^\circ$) reflections, are: $a = 18.234(2) \text{ \AA}$, $b = 8.816(1) \text{ \AA}$, $c = 5.192(1) \text{ \AA}$. Unit-cell parameters for the two clinopyroxene crystals (space group *C2/c*), determined by centering the same set of 25 high- θ ($20\text{--}28^\circ$) reflections, are: $a = 9.714(1) \text{ \AA}$, $b = 8.916(1) \text{ \AA}$, $c = 5.252(1) \text{ \AA}$, $\beta = 105.93(1)^\circ$, $V = 437.4(2) \text{ \AA}^3$ and $a = 9.725(1) \text{ \AA}$, $b = 8.915(1) \text{ \AA}$, $c = 5.258(1) \text{ \AA}$, $\beta = 106.02(1)^\circ$, $V = 438.2(2) \text{ \AA}^3$ for the crystals from 1006 and 927 run product, respectively. Several grains of KAlSi_3O_8 were hand-picked from the run product 1024, finely ground in an agate mortar, and prepared as a thin layer on a silicon single-crystal. Step-scan powder diffraction data, with 0.02° as step and 9 s as counting time, were recorded with an automated diffractometer (Philips) in the 2θ range $2\text{--}70^\circ$ using Cu $K\alpha$ radiation and silicon powder (NBS 640a) as external standard. Detailed examination of the observed pattern shows the absence of diffraction peaks belonging to any crystalline phase other than high sanidine.

The Raman analyses of glasses in order to verify the presence of water were performed using the RM-1000 micro-Raman spectrometer. The samples were exited at

room temperature with the 514-nm line of Ar-laser. The magnification of the objective was $50\times$.

Experimental results

Textural relationships in the run products

The run conditions and results of each experiment are presented in Table 1. Textures in the experiments produced at $1,300\text{--}1,450^\circ\text{C}$ at the $\text{CaMgSi}_2\text{O}_6$ content in the system $>70 \text{ mol}\%$ are ambiguous because of spectacular dendritic clinopyroxene crystals (Fig. 1a), which are interpreted as quenched crystals. Orthopyroxene in the experiments 929 and 1159 (Table 1) forms large (up to $50 \mu\text{m}$) crystals, which locally contain minute clinopyroxene inclusions (Fig. 1b). In the $\text{CaMgSi}_2\text{O}_6$ -rich portion of the join, clinopyroxene forms elongated prismatic crystals up to $100 \mu\text{m}$ in length (Fig. 1c, d). In the KAlSi_3O_8 -rich portion of the join, clinopyroxene forms both clusters of small crystals and relatively large anhedral crystals in the products of quenching (Fig. 1e). At the solidus, clinopyroxene forms acicular crystals (Fig. 1f). As a rule, garnet grains are equant (Fig. 1d). In the experiments produced in the diopside-rich portion of the join at temperatures $1,000\text{--}1,100^\circ\text{C}$, garnet form large (up to $100 \mu\text{m}$ in diameter) euhedral crystals and locally includes clinopyroxene. Sanidine forms large subhedral crystals (up to $100\text{--}300 \mu\text{m}$) containing clinopyroxene inclusions (Fig. 1e). In the high-temperature experiments, sanidine crystals are surrounded by quench zones and contain “quench pockets”, which are composed of acicular sanidine intermixed with residual glass (Fig. 1e). At lower temperatures, sanidine crystallizes as euhedral grains (Fig. 1f).

All experiments contain quenched aluminosilicate melt. In the experiments conducted at temperatures above $1,200^\circ\text{C}$ at 6 GPa and above $1,000^\circ\text{C}$ at 3.5 GPa, the quenched melt appears as relatively homogeneous bubble-free glasses resulting from rapid quenching (about $300\text{--}400^\circ\text{C/s}$). Locally, there are aggregates of acicular crystals, irregularly dispersed in the homogeneous glass (Fig. 1c, d). Glass is extremely rare in the experiments near the inferred solidus. In most cases, the products of the melt quenching near the solidus are aggregates of acicular crystals.

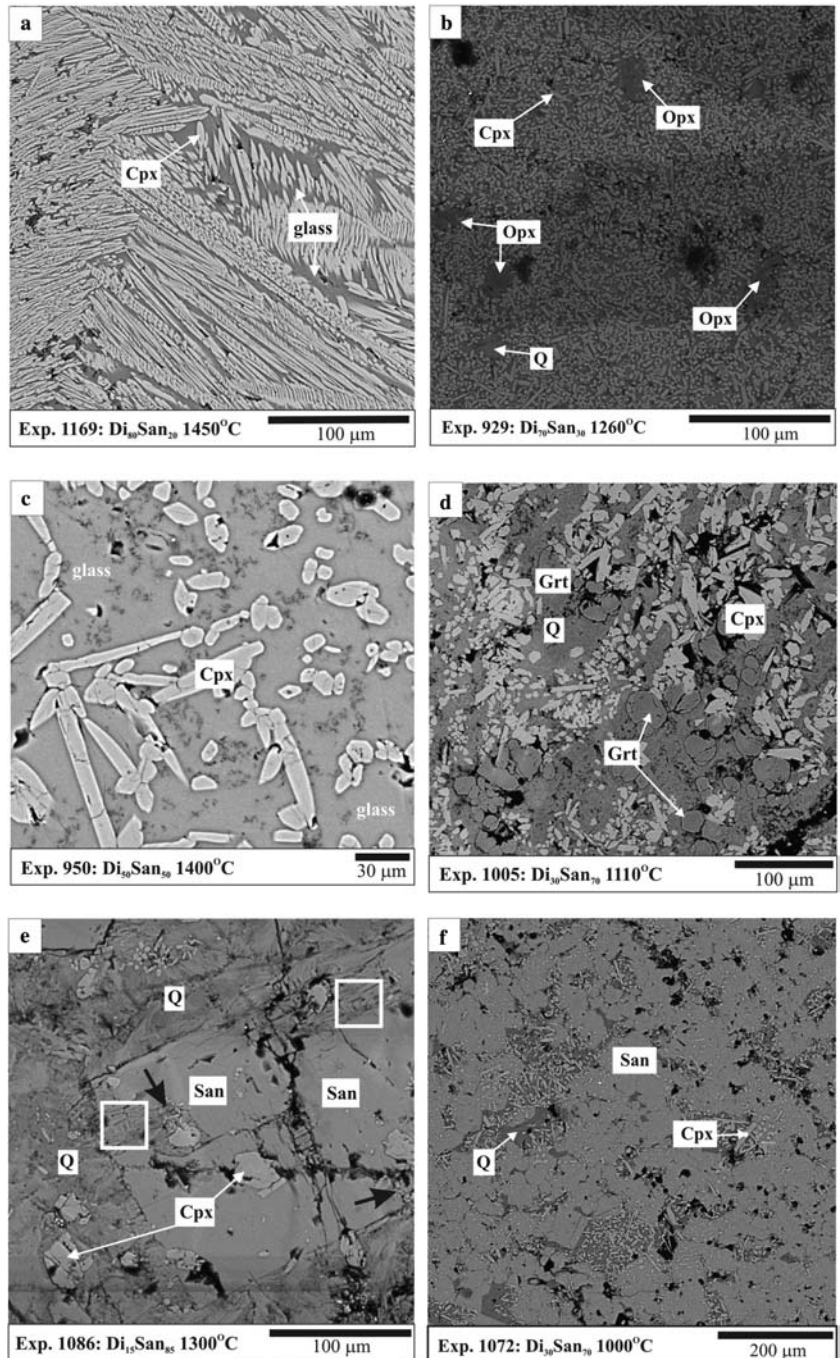
Specific zoned distribution of phase assemblages is characteristic for the experiments containing Grt, Cpx, San, and quenched melt (experiments 1089, 1090, 1121; Table 1). For example, experiment 1090 shows three zones (from bottom to top of Fig. 2a): (1) the Cpx + Grt + glass zone, (2) the glassy zone, and (3) the Cpx + San + Q zone. The zone (1) is an aggregate of subhedral clinopyroxene grains intergrown with garnet (Fig. 2b). The Cpx + San + Q zone is composed of subhedral San and Cpx grains with “pockets” of quenching textures (Fig. 2c). Textures of both zones resemble a cumulate of crystals in the melt. This appearance would imply a

gravitational separation of phases during long-time runs (4–6 h): phases of higher density (Grt and Cpx) settle onto the capsule bottom, whereas lower-density sanidine tends to float up. Similar Cpx and Grt aggregation on the capsule bottom is observed in some San-free experiments, as well.

Partial replacement of large sanidine crystals with aggregates Cpx + quench is identified in the experiments 1046, 1086, 1087, and 1194 (Table 1; Figs. 1e and 2d). This texture would imply the schematic reaction $\text{San} + \text{melt} \rightarrow \text{Cpx}$ in progress. The formation of the texture is presumably related to kinetics of phase crystalliza-

tion in the sanidine-rich portion of the join. Large euhedral sanidine crystals (up to 300 μm) in this portion of the join (Fig. 1e) may have been grown immediately after pressurization but in the course of heating (see the experimental procedure above); growing sanidine crystals captured separate diopside grains, as well. At the run conditions, crystalline sanidine appeared to be as the starting mineral. In the course of the experiment, sanidine was re-equilibrated by means of interaction with melt to form new clinopyroxene. Diopside shielded by sanidine does not display any reaction (Fig. 1e).

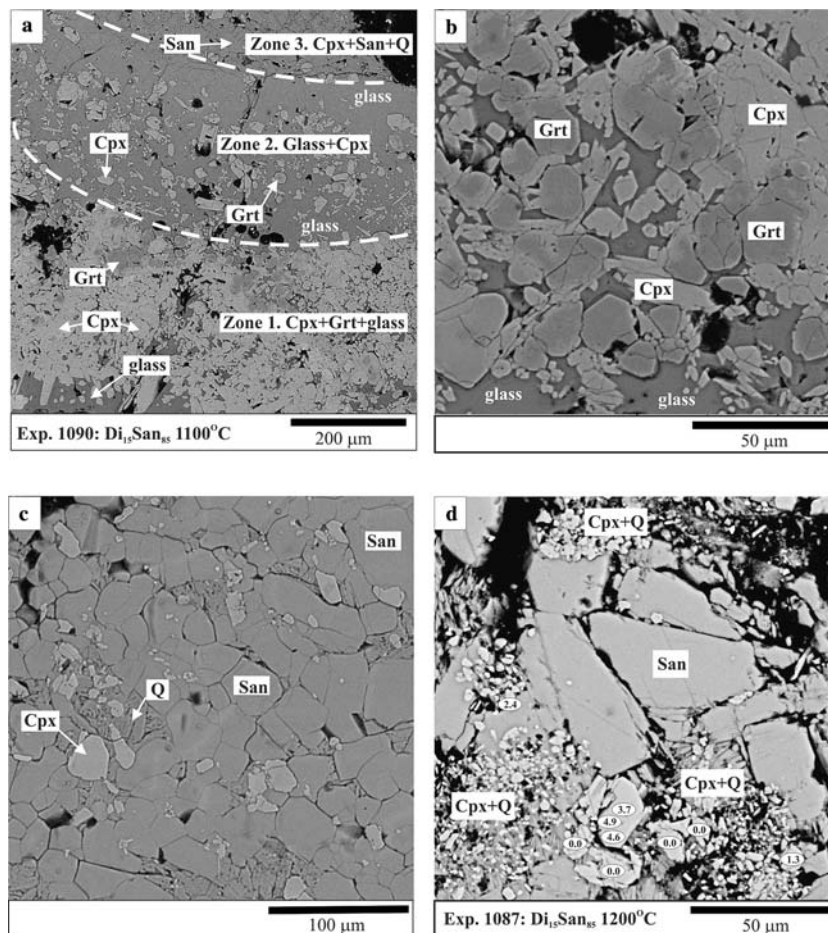
Fig. 1 Textural phase relationships in the run samples synthesized in the join $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$ at 6 and 3.5 GPa. **Plate a** Skeleton crystals of Cpx and SiO_2 -rich glass. These textures create difficulties in identification of the liquidus in the $\text{CaMgSi}_2\text{O}_6$ -rich portion of the join. **Plate b** Large euhedral Opx crystals and small Cpx crystals plunged into the quenching products (Q). **Plate c** Euhedral Cpx crystals in the aluminosilicate melt at the liquidus of the join. Darker areas in the glass are composed of acicular phases, the quenching products (Q). **Plate d** Crystals of Cpx and Grt in the matrix that is composed of quenched melt (Q). The quenched melt around some crystals is presented by homogeneous glass. **Plate e** Large San crystals and potassium-poor Cpx crystals shielded by the sanidine crystals. Quenching zones and “pockets” at the San crystals (indicated with *squares*) are composed of acicular aggregates of sanidine and residual glass. *Black arrows* mark the places, where sanidine crystals are partially replaced with the aggregates of K-bearing Cpx and quench. **Plate f** Euhedral sanidine and acicular clinopyroxene crystals at the solidus of the join $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$ at 3.5 GPa. Sanidine grains contain numerous inclusions of clinopyroxene, giving an appearance of the eutectic-like texture; *gray* areas between sanidine crystal are quenching products (Q)



Phase composition

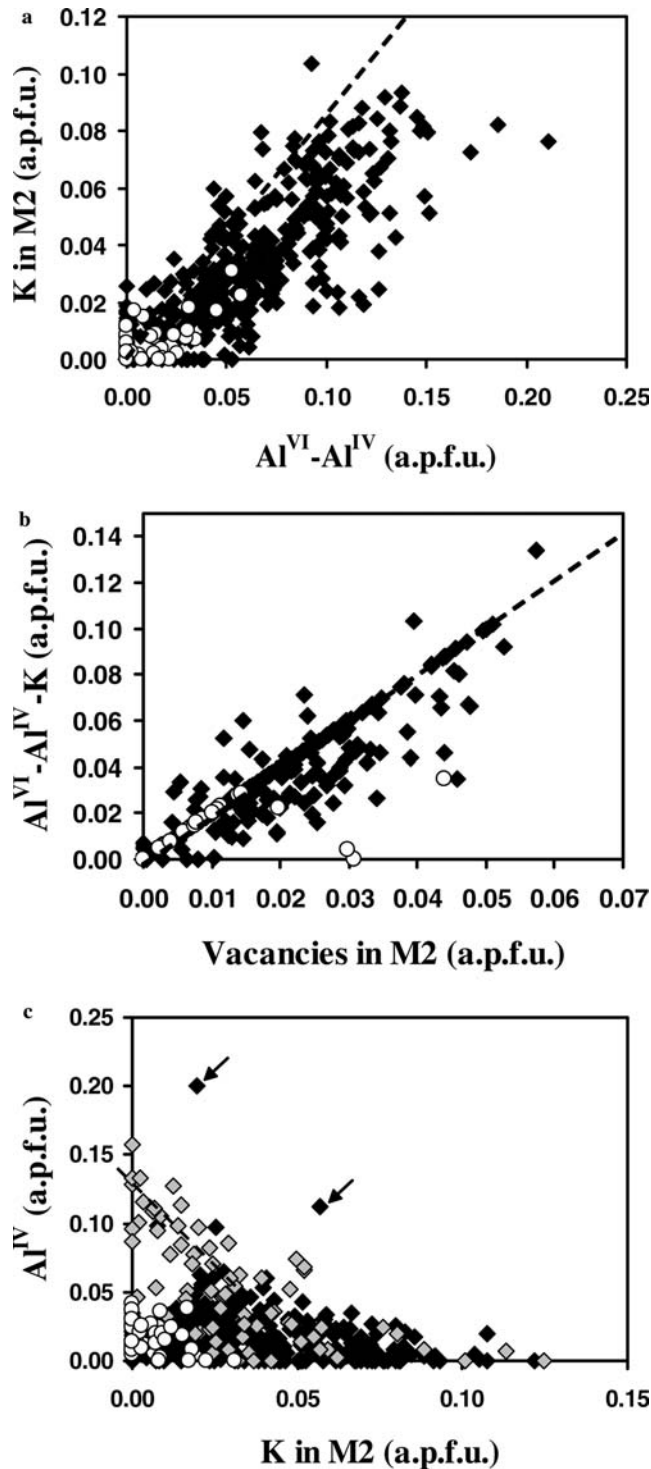
The mean grossular content of *garnets* produced at 6 GPa varies from 20 to 35 mol% (Table 2). Garnets show either concentric (boundaries of zones are parallel to the crystal outlines) or spotty (Ca-rich zones form spots in the grains or are attached to the contacts with clinopyroxene) zoning. Grossular content in garnets from the experiment 1090 varies from 40 up to 55 mol%

Fig. 2 Specific textural features of the sanidine-bearing samples synthesized in the join $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$ at 6 GPa. **Plate a** Zoned distribution of phase assemblages produced by separation of phases by gravity. Three zones are distinguished in the sample (from the bottom to the top): Zone 1—Cpx + Grt + glass; Zone 2—Cpx + glass (with rare Grt); and Zone 3—Cpx + San + quench. *Dashed lines* roughly mark boundaries between the zones. Relationships of phase in the first and the third zones are shown in plates 2b and 2c. These plates are outside of the plate 2a. **Plate b** Detailed view of the cumulate-like aggregates and intergrowths of Cpx and Grt with minor interstitial glass in the zone 1 in the Plate 2a. **Plate c** Detailed view of subhedral San and Cpx grains with “pockets” of quenching textures (Q) in the zone 3 in the Plate 2a; quenching textures are composed of acicular sanidine and residual glass. **Plate d** Replacement of San crystals with aggregates of Cpx + Q reflecting the assumed reaction $\text{San} = \text{KJd} (\text{in Cpx}) + \text{L}$ in the KAlSi_3O_8 -rich portion of the join. Numbers in the Cpx fields show the KJd content in this clinopyroxene (mol%)



(Table 2). Centers of grains show lower N_{Ca} , while rims contacting to clinopyroxene are affected by a local increase of the grossular content (Fig. 1e). The zoning of garnets in this experiment (as well as in 1089), probably, resulted from re-equilibration of composition after settling of garnet on the capsule bottom during the run. No systematic variation of garnet composition with temperature is found, but garnets in experiments at the lowest temperatures (978 and 1037; Tables 1 and 2) are very Ca-rich ($N_{\text{Ca}} = 80\text{--}90$ mol%).

Clinopyroxenes synthesized at 6 GPa show the positive correlation of the K content with $\text{Al}^{\text{M1}}\text{-Al}^{\text{IV}}$ (Fig. 3a), which reflects the substitution $\text{Mg}^{\text{M1}} + \text{Ca}^{\text{M2}} \Leftrightarrow \text{Al}^{\text{M1}} + \text{K}^{\text{M2}}$ (e.g., Harlow 1997; Chudinovskikh et al. 2001; Safonov et al. 2002, 2003) and the presence of the KJd end-member in the clinopyroxenes. The $\text{K} \Leftrightarrow \text{Ca}$ substitution into the M2 site of the clinopyroxene structure was also verified by the X-ray single-crystal refinements of two compositionally homogeneous crystal fragments from the experiments 1006 and 927, containing 0.60 and 0.81 wt% of K_2O , respectively. A notable excess of Al in the M1 site with respect to Ca-Ts and KJd (i.e., $\text{Al}^{\text{M1}}\text{-Al}^{\text{IV}}\text{-K}$) (Fig. 3a) is directly correlated as 2 : 1 with vacancies in the M2 site ($\square^{\text{M2}} = 1 - \text{Ca} - \text{K} - \text{Mg}^{\text{M2}}$) in most of clinopyroxenes (Fig. 3b). This correlation is evidence for the Ca-Esk end-member in KCpx. In the most cases, the Al^{IV}



content (i.e. Ca-Ts end-member) of clinopyroxenes does not exceed 0.05 a.p.f.u. (Table 2; Fig. 3c). The highest concentrations of Al^{IV} up to 0.10–0.19 a.p.f.u. are characteristic for some cores of large Cpx crystals, Cpx inclusions in garnets, as well as Cpx intergrown with garnet in the sanidine-bearing experiments 1089, 1090, and 1121 (Table 2; Fig. 3c). For the Ca-Ts-rich clinopyroxenes, Al^{IV} shows strong negative correlation with

Fig. 3 General compositional features of clinopyroxenes synthesized in the join CaMgSi₂O₆-KAlSi₃O₈ at 6 GPa (*black diamonds*) and 3.5 GPa (*white circles*). **a** Correlation of K in the M2 site with Al in the M1 site (excluding Al bound to the CaAl₂SiO₆ end-member). *Dashed line* shows a theoretical relation of these parameters corresponding to the two-site heterovalent isomorphism $Mg^{M1} + Ca^{M2} \Leftrightarrow Al^{M1} + K^{M2}$. **b** Correlation of vacancies in the M2 site with Al in the M1 site (excluding Al bound to the CaAl₂SiO₆ and KAlSi₂O₆ end-members). *Dashed line* shows a theoretical relation of these parameters corresponding to the two-site heterovalent isomorphism $Mg^{M1} + 0.5Ca^{M2} \Leftrightarrow Al^{M1} + 0.5\Box^{M2}$. **c** Variation of Al^{IV} and its negative correlation with K in the M2 site for clinopyroxenes enriched in Ca-Ts (*dashed line*). In this diagram, *black diamonds* show clinopyroxenes from the sanidine-free samples, while *gray diamonds* represent clinopyroxenes from the sanidine-bearing samples produced at 6 GPa. *Arrows* point to the clinopyroxenes included in garnets

KJd (Fig. 3c). All clinopyroxenes contain Mg in the M2 site, i.e., CEn end-member (Table 2). Similar to the Ca-Ts end-member, highest CEn content is observed in cores of large clinopyroxene crystals and in clinopyroxenes included in garnets.

Thus, clinopyroxenes produced in the system CaMgSi₂O₆-KAlSi₃O₈ at 6 GPa are complex solid solutions of Di, KJd, Ca-Esk, CEn, and Ca-Ts. The KJd content of clinopyroxenes in the sanidine-free experiments produced at $N_{San}^{system} = 40\text{--}70$ mol% regularly increases upon decreasing temperature and reaches 10 mol% at $\sim 1,000^\circ\text{C}$ (Fig. 4). In contrast, the CEn content regularly decreases upon decreasing temperature. The Ca-Ts content tends to decrease slightly with decreasing temperature, whereas the Ca-Esk in clinopyroxenes is approximately constant at temperatures above 1,000°C.

The most interesting variations of composition are observed in Cpx in the sanidine-bearing experiments. The high temperature garnet-free experiment 1046 contains clinopyroxene with a mean KJd content about 5.7 mol% (Table 2), which is slightly higher than the KJd content in Cpx from the sanidine-free experiments at similar temperature (Fig. 4). In the experiments 1086, 1087, 1088, and 1194 (Table 1), large sanidine crystals include clinopyroxenes containing 0.0–0.4 wt% of K₂O, but no alumina to compensate potassium (Fig. 1e and Table 2). The irregular presence of uncompensated potassium in these clinopyroxene, probably, corresponds to K-fluorescence from sanidine and implies an absence of the KJd end-member. However, in addition to these clinopyroxenes (actually, diopside), experiments 1087 and 1194 (Table 1 and 2) contain relatively large clinopyroxene crystals enriched in KJd (4–8 mol%), Ca-Ts (2–7 mol%), and Ca-Esk (4–9 mol%). The detailed examination shows that KCpx grains are always coexist with the quenched melt, whereas diopside crystals are mostly associated with sanidine aggregates. Partially decomposed diopside grains are locally situated very close to subhedral KCpx grains, while aggregates of KCpx grains and quenched melt replace sanidine along grain boundaries (Fig. 2d). The zoned run samples 1121 and 1090 (Table 1) contain diopside as well, which is

surrounded by the sanidine aggregates (Fig. 2a, c). In addition to diopside, these samples contain Al-rich clinopyroxene in the Grt–Cpx cumulates and as separate crystals in glass (Fig. 1e). The Ca-Ts content varies within the range 4–7 mol%, but increases up to 15 mol% in contacts with garnet. The increase of the Ca-Ts content results in appreciable decrease of the KJd and Ca-Esk. Therefore, most of clinopyroxenes from the experiments 1090 and 1121 deviate from the general dependence of the KJd content on temperature (Fig. 4). Nevertheless, the concentration of the KJd component (about 5 mol%) in relatively low-Ca-Ts KCpx is in a good agreement with the general dependence in Fig. 4. Further reduction in temperature results in an increase of the KJd in clinopyroxene (e.g., experiments 1024 and 1136). The estimated KJd concentration in some clinopyroxenes from the experiments 1022 and 978 produced at temperatures 940 and 890°C (Table 1), respectively, reaches ~12 mol% (Table 2), which is consistent with the general dependence of the KJd content in KCpx on temperature, as well (Fig. 4). However, formulae of clinopyroxene from the experiment 1022 show a notable excess of Si over 2.0 a.p.f.u., probably, because of contamination with SWd. Therefore, its end-members cannot be estimated properly.

Thus, compositions of KCpx from the sanidine-bearing assemblages show a good correlation with temperature (Fig. 4). The deviations of the clinopyroxene compositions from this dependence can be explained by the significant textural inhomogeneity of the sanidine-bearing run samples.

Clinopyroxenes synthesized at 3.5 GPa show considerably lower KJd content (Table 2; Fig. 3a). Nevertheless, the potassium concentration in these clinopyroxenes increases with decreasing temperature as well. Just the run products 1072, obtained near the solidus (Table 1 and Fig. 1g), contain clinopyroxene with 0.6 wt% K₂O. However, we suggest that this apparent K₂O concentration is mostly related to fluorescence from sanidine during microprobe analyses of these acicular crystals (Fig. 2). Clinopyroxenes produced at 3.5 GPa contain Ca-Esk, as well, but the concentration does not exceed 3 mol%.

Orthopyroxene from both the experiments 1159 and 929 contain up to 1.5 wt% of Al₂O₃ and CaO 1.2–2.5 wt% of CaO, suggesting low concentrations of Mg-Ts and Wol end-members (Table 2).

Sanidine composition in all experiments obeys the stoichiometry of KAlSi₃O₈. The total content of CaO and MgO content does not exceed 0.5 wt% (Table 2).

Si-wadeite in experiment 1022 (Tables 1 and 2) contains CaO, MgO, and Al₂O₃ with reduction in K₂O, which is interpreted as minute inclusions of clinopyroxene, garnet, and sanidine contaminants.

Glasses in all run products are characterized by relatively high SiO₂ (55–75 wt%), Al₂O₃ (11–19 wt%), and K₂O (12–22 wt%) contents at CaO + MgO concentration less than 7 wt%. Composition of glasses produced at 6 GPa varies with temperature and bulk

composition. For glasses synthesized in the diopside-rich portion of the join, the SiO₂ concentration is nearly constant, whereas it increases for glasses synthesized in the sanidine-rich portion. The join K₂O slightly decreases with increasing temperature for all glasses. The Al₂O₃ content in glasses is approximately constant over the temperature range of experiments.

Using Raman spectroscopy, we verified the presence of water in glasses from experiments 950, 1011, and 1109 (Table 1). In addition to peaks in the low-frequency range corresponding to vibration of tetrahedral cations, spectra showed a wide band with a peak in the frequency range 3,500–3,550 cm⁻¹ attributed to stretching vibrations from molecular water (e.g., Holz

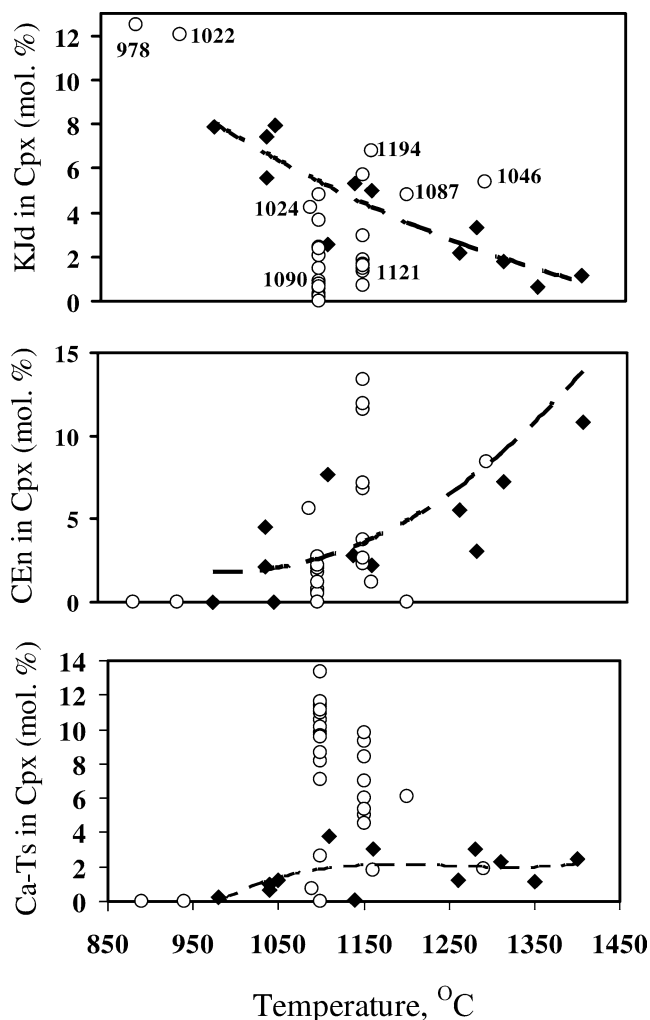


Fig. 4 Temperature dependence of the mean KJd, CEn, and Ca-Ts contents for clinopyroxenes from sanidine-free samples (*black diamonds*) and sanidine-bearing samples (*white circles*) produced at 6 GPa. Labels indicate the sanidine-bearing run samples discussed in the text. For the experiments 1090 and 1121, all measured clinopyroxenes are plotted in order to demonstrate a variability of the clinopyroxene composition in these samples (see text). *Dashed lines* represent fits for clinopyroxene compositions from sanidine-free samples

et al. 1996). No bands of OH-groups, which would be expected at relatively high water content, were present in the spectra. Analyses of glasses usually show a total oxide concentration close to 100 wt% (Table 2), which implies low water concentration in these glasses, as well. Glass is very rare in the experiments produced at temperatures below 1,100°C, whereas products of the melt quenching are presented by aggregates of acicular crystals, probably, of mica-like phases. Analyses of these quenching aggregates show the total oxide concentration 95–97 wt% (Table 2). Their formation is related to local accumulation of water (since neither Cpx nor Grt and San contains water) in the residual melt. The presence of these quenching textures in the low-temperature experiments makes locating the solidus in T–X space difficult.

T–X diagrams for the join CaMgSi₂O₆–KAlSi₃O₈ at 6 and 3.5 GPa

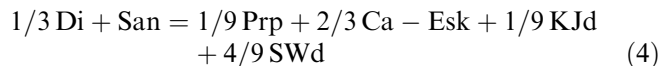
Topology and T–X diagram for the join CaMgSi₂O₆–KAlSi₃O₈ at 6 GPa

The system Di–San at ambient pressure is a pseudobinary join within the system CaMgSi₂O₆–KAlSi₂O₆–SiO₂, characterized by crystallization of Lc at liquidus and peritectic reaction Lc + L = San (Schairer and Bowen 1938). At pressures above 2 GPa, sanidine melts congruently (e.g., Lindsley 1966) up to about 6.5 GPa. Sanidine is unstable at higher pressure with the formation of the assemblage SWd + Ky + Cos (Kinomura et al. 1975; Urakawa et al. 1994; Yagi et al. 1994; Akaogi et al. 2004). But diopside is stable and melts congruently up to pressure ~16 GPa (Gasparik 1996). Thus, our experiments (Table 1) were conducted within the stability fields of both the San and Di up to temperatures of their congruent melting.

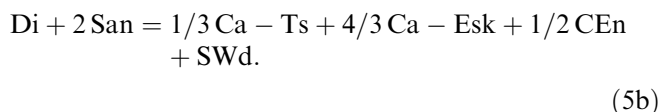
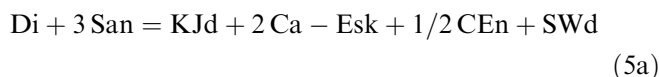
Phase relations on the join CaMgSi₂O₆–KAlSi₃O₈ can be described in terms of the five-component system K₂O–CaO–MgO–Al₂O₃–SiO₂. As in our previous study (Safonov et al. 2003), we consider the four-component system K₂O–(Ca, MgO)–Al₂O₃–SiO₂ in order to simplify the topological analysis of the join. This simplification assumes that the Ca/Mg in garnets and clinopyroxenes is constant, while clinopyroxene is the only potassium-bearing solid solution. The true Ca/Mg is not constant in the crystallized phases (see below). However, this simplification allows derivation of a general topology of the join CaMgSi₂O₆–KAlSi₃O₈.

Phases at the solidus on the join Di–San at 6 GPa are Cpx_{ss} (the Di–KJd–CEn–Ca–Esk–Ca–Ts solid solution), Grt_{ss} (the Prp–Grs solid solution), and San. SWd was observed with Cpx_{ss} + San in experiments at $N_{\text{San}}^{\text{system}} = 95 \text{ mol}\%$ (see experiment 1022 in Tables 1 and 2). Opx_{ss} (the En–Mg–Ts–Wol solid solution) is identified in the CaMgSi₂O₆-rich portion of the system (Table 1), but its presence at the solidus is not

confirmed. The Cpx_{ss}, Grt_{ss}, and SWd are products of reactions between starting Di and San:

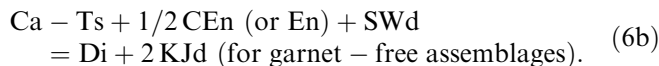
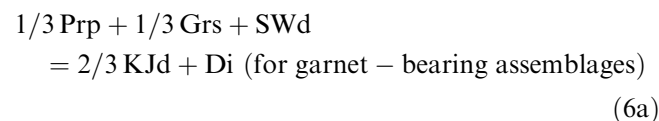


Compositions of Grt, KCpx, and SWd are located off the binary join Di–San, i.e., their compositions cannot be expressed as linear combination of the two end-member components (Di and San). This feature allows classification of the join as pseudobinary (Ehlers 1972; Maaløe 1985; Litvin 1991). The positions of Di, Grt, SWd, San compositions within the K₂O–(Ca, MgO)–Al₂O₃–SiO₂ tetrahedron are shown in Fig. 5a. These end-members define a smaller tetrahedron, one edge of which is the join Di–San. In addition, clinopyroxene is a complex solid solution of Di, KJd, Ca–Esk, CEn, and Ca–Ts, whose formation also results from reactions between the starting Di and San:



At relatively low temperatures, the CEn end-member exsolves from the Cpx_{ss} to form a separate phase, i.e., the Opx_{ss} (see experiments 929 and 1159 in Table 1 and Fig. 2a).

Reactions (4) and (5) contain SWd as a product phase. However, we surely identified SWd only at $N_{\text{San}}^{\text{system}} = 95 \text{ mol}\%$ (Table 1), but this phase was not observed in the dominant portion of the interpreted solidus. Si-wadeite could be unobserved because of its low content. However, since no traces of SWd were found during the special examination (repeated polishing and the BSE mapping), we assume that SWd is not stable in this portion of the join, because of its consumption to form additional KJd in the clinopyroxene solid solution:



In the San-rich portion of the join, the bulk composition is suitable for coexistence of SWd with KCpx and San.

Due to the formation of Cpx_{ss} and Opx_{ss}, the large tetrahedron Di–Grt–SWd–San splits into pseudo-tetrahedra (Cpx + Opx)_{ss}, (Cpx + Opx)_{ss}–Grt, (Cpx + Opx)_{ss}–Grt–San, (Cpx + Opx)_{ss}–SWd, (Cpx + Opx)_{ss}–

Gr_t-SW_d, and (Cpx + Opx)_{SS}-Gr_t-SW_d-San (Fig. 5b), which correspond to subsolidus assemblages within the system Di-Gr_t-SW_d-San. Similar to the join Di-Lc at 7 GPa (Safonov et al. 2003), owing to both the Cpx_{SS} and the Opx_{SS} formation, the beginning of a true join is slightly displaced inward of the (Cpx + Opx)_{SS} tetrahedron. In addition, because of SW_d formation, the end of the join is slightly shifted inward the large tetrahedron (Cpx + Opx)_{SS}-Gr_t-SW_d-San, but is located very close to the San apex (the end of the join is further denoted as San*). Therefore, the Di-San serves as a fictive join. Displacement of the boundaries of the true join is another evidence of its pseudobinary character (Ehlers 1972; Maaløe 1985; Litvin 1991). Starting from some point inside the (Cpx + Opx)_{SS} tetrahedron, the true Cpx_{SS}-San* join penetrates four tetrahedra (Fig. 5b): (Cpx + Opx)_{SS}, (Cpx + Opx)_{SS}-Gr_t, (Cpx + Opx)_{SS}-Gr_t-San, and (Cpx + Opx)_{SS}-Gr_t-SW_d-San.

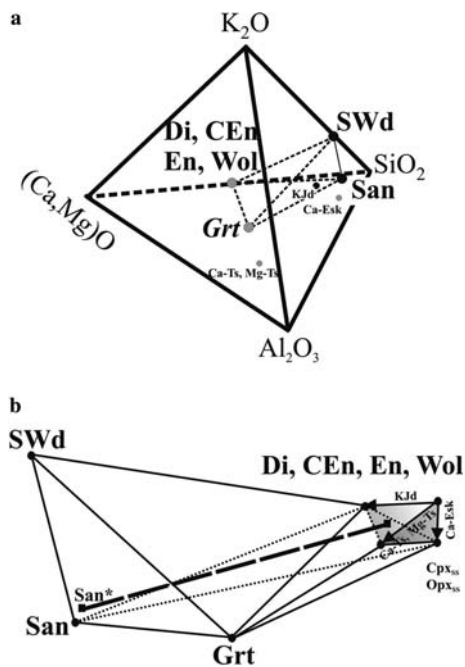


Fig. 5 Topology of phase relations in the join $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$ at 6 GPa (see explanation in the text). **a** Position of phases and end-members forming in the join $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$ in the $\text{K}_2\text{O}\text{-(Ca, Mg)O-Al}_2\text{O}_3\text{-SiO}_2$ tetrahedron. Solid lines show visible tie-lines, while dashed lines indicate invisible tie-lines. Black data points are within the $\text{K}_2\text{O-Al}_2\text{O}_3\text{-SiO}_2$ plane. Gray data points are within the $\text{(Ca, Mg)O-Al}_2\text{O}_3\text{-SiO}_2$ plane. **b** Positions of the $(\text{Cpx} + \text{Opx})_{\text{SS}}\text{-San}^*$ join (thick dashed line) with respect to phases and end-members in the system $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$. Gray tetrahedrons denote the Cpx and Opx solid solutions: Di, CEn, KJd, Ca-Esk, Ca-Ts and En, Mg-Ts, Wol, respectively. Arrows are directed from Di (En) to other end-members of the Cpx_{SS} and Opx_{SS}, whereas black points at the arrow ends schematically illustrate a maximal solubility of corresponding end-members in the Cpx_{SS}. Black squares schematically show a displacement of the beginning and the end of the true join $(\text{Cpx} + \text{Opx})_{\text{SS}}\text{-San}^*$ inward both the $(\text{Cpx} + \text{Opx})_{\text{SS}}$ and $(\text{Cpx} + \text{Opx})_{\text{SS}}\text{-Gr}_t\text{-SW}_d\text{-San}$ tetrahedrons. Solid lines show visible tie-lines, while dashed lines indicate invisible (back-edge) tie-lines

The end of the join is located inside the $(\text{Cpx} + \text{Opx})_{\text{SS}}\text{-Gr}_t\text{-SW}_d\text{-San}$ tetrahedron close to the San point. This position of the join $\text{Cpx}_{\text{SS}}\text{-San}^*$ corresponds to the following phase assemblages in the subsolidus: $(\text{Cpx} + \text{Opx})_{\text{SS}}$, $(\text{Cpx} + \text{Opx})_{\text{SS}}\text{-Gr}_t$, $\text{Cpx}_{\text{SS}}\text{-Gr}_t\text{-San}$, and $\text{Cpx}_{\text{SS}}\text{-Gr}_t\text{-SW}_d\text{-San}$. Presence of Opx_{SS} implies that the subsolidus region of at $N_{\text{San}}^{\text{system}} < 10 \text{ mol}\%$ could split into free additional fields: Cpx_{SS} , $\text{Cpx}_{\text{SS}} + \text{Opx}_{\text{SS}}$, and $\text{Cpx}_{\text{SS}} + \text{Opx}_{\text{SS}} + \text{Gr}_t$. In high-temperature experiments, either Cpx (in the diopside-rich portion of the join; Fig. 1c) or San (in sanidine-rich portion of the join) are found in glass. That implies that the join includes two liquids only: Cpx_{SS} and San. Opx_{SS} was observed only at 1,200–1,260°C in the run products at $N_{\text{San}}^{\text{system}} < 30 \text{ mol}\%$ (runs 929 and 1159; Table 1), while it was not detected in the solidus at $N_{\text{San}}^{\text{system}} > 10 \text{ mol}\%$. This implies that Opx disappears at cooling, probably, via peritectic reaction to form the assemblage Cpx + Gr_t, so the Opx_{SS} could be present in solidus only at $N_{\text{San}}^{\text{system}} < 10 \text{ mol}\%$. The assumed peritectic is very similar to that in the join Di-Prp at 3 GPa (O'Hara 1963), but occurs at lower temperature owing, probably, to the presence of the K-Si-rich melt. In the absence of Opx at higher temperature, the CEn end-member dissolves in Cpx_{SS} (Fig. 4).

Thus, we established a possible sequence of phase fields in the subsolidus and implied liquidus relations in the join $\text{Cpx}_{\text{SS}}\text{-San}^*$ at 6 GPa. This allows construction of the $T\text{-X}$ diagram using the data from Table 1 and the rules on construction of the diagrams for the pseudobinary systems (Rhines 1956). The pseudobinary join $\text{Cpx}_{\text{SS}}\text{-San}^*$, which is a chemical analogue of the system $\text{CaMgSi}_2\text{O}_6\text{-KAlSi}_3\text{O}_8$, is characterized by the following features (Fig. 6).

1. Up to $N_{\text{San}}^{\text{system}} = 85 \text{ mol}\%$, Cpx_{SS} is a liquidus phase. Only at the higher $N_{\text{San}}^{\text{system}}$ sanidine does become a liquidus phase.
2. The assemblage $\text{Opx}_{\text{SS}} + \text{Cpx}_{\text{SS}} + \text{L}$ is stable in the $\text{CaMgSi}_2\text{O}_6$ -rich portion of the join. Peritectic reaction $\text{Opx}_{\text{SS}} + \text{Cpx}_{\text{SS}} + \text{L} = \text{Cpx}_{\text{SS}} + \text{Gr}_t$ is proposed to explain the absence of Opx in the solidus at the $N_{\text{San}}^{\text{system}} > 10 \text{ mol}\%$.
3. The join $\text{Cpx}_{\text{SS}}\text{-San}^*$ is characterized by a wide field of the joint crystallization of Cpx and pyrope-grossular Gr_t, which corresponds to the $\text{Cpx}_{\text{SS}} + \text{Gr}_t + \text{L}$ cotectic.
4. Below the liquidus, sanidine is stable only at $N_{\text{San}}^{\text{system}} > 70 \text{ mol}\%$. The cotectic assemblages $\text{Cpx}_{\text{SS}} + \text{San} + \text{L}$, $\text{Cpx}_{\text{SS}} + \text{Gr}_t + \text{San} + \text{L}$, $\text{Cpx}_{\text{SS}} + \text{San} + \text{SW}_d + \text{L}$ are stable in the sanidine-rich portion of the join.
5. The eutectic $\text{Cpx}_{\text{SS}} + \text{Gr}_t + \text{San} + \text{SW}_d + \text{L}$ is situated very close to the San apex. The position of the entire solidus is unconstrained, because ambiguous interpretation of quenching textures related to the local accumulation of water in the low-temperature experiments. Position of the solidus line in Fig. 6 is conventional.

The phase diagram in Fig. 6 has to be considered as preliminary, because (1) the $\text{CaMgSi}_2\text{O}_6$ -rich portion of the diagram is not well experimentally constrained and requires further experiments, especially for the Opx – Cpx – Grt relationships at the assumed peritectic and the solidus; (2) exact position of the solidus is not established. In addition, no variation of Ca-content of Grt and the CEn, Ca-Esk and Ca-Ts components in clinopyroxene are accounted for. These features cannot be clearly indicated on the 2D diagram for the multi-component system. Experimental data allow confirmation of the general topology of the join and specification of major phase fields, assumed from the topological analysis.

Topology and the T–X diagram for the join $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 at 3.5 GPa

Grt and SWd are absent on the join Di–San at 3.5 GPa (Table 1). However, Cpx at this pressure also contains low concentrations of KJd, Ca-Esk, and CEn (Table 2; Fig. 3a, b). The introduction of these end-members in the Cpx_{ss} results in the displacement of Cpx from diopside and, subsequently, from the Di–San tie-line. This feature corresponds to a slight deviation of the Cpx_{ss} –San join at 3.5 GPa from a truly binary join (Fig. 7). The topology of the join at 3.5 GPa is similar to the topology of the system Di–San– H_2O at 1 GPa of water pressure characterized by the formation of the Di–Ca–Ts solid solution (Yoder and Upton 1971). Clinopyroxene is a liquidus phase at $N_{\text{San}}^{\text{system}} < 70$ mol%, whereas San crystallizes at liquidus at the higher $N_{\text{San}}^{\text{system}}$. Ternary eutectic $\text{Cpx} + \text{San} + \text{L}$ occurs between 1,000 and 1,050°C.

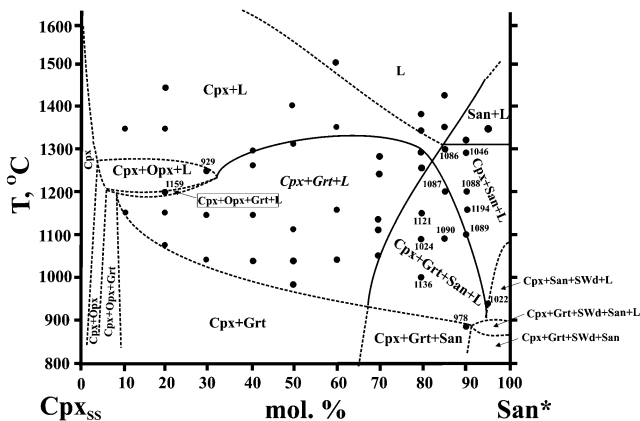


Fig. 6 Preliminary phase T–X diagram for the system $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 , i.e., join Cpx_{ss} – San^* at 6 GPa. Solid lines denote experimentally established boundaries of phase fields; dashed lines indicate the assumed boundaries of phase fields; labeled data points relate to the specific experiments (Table 1) discussed in the text

Discussion and application to the natural assemblages

Experiments at 6 and 3.5 GPa demonstrate that the K-rich melts in the system $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 could produce common mantle mineral assemblages, such as $\text{Cpx} + \text{Grt}$, $\text{Cpx} + \text{Opx}$, $\text{Cpx} + \text{San}$, or $\text{Cpx} + \text{Grt} + \text{San}$. At 6 GPa, the assemblage $\text{Grt} + \text{Cpx}$ coexists with K–Al–Si-rich melt down to about of 1,000–1,100°C (Fig. 6). These temperatures are appreciably lower than temperatures of crystallization (at least above 1,300°C) for the assemblage $\text{Grt} + \text{Cpx}$ from potassium-free melts, including Na and Fe-bearing systems (see Litvin 1991 and references therein for a review of the phase diagrams in model and natural eclogite systems). Even if we take into account the presence of the uncontrolled amount of water in the starting gels, the lowering of liquidus and solidus of sanidine-bearing system in comparison to the potassium-free systems is evident. Low solidus temperatures are caused probably by the high SiO_2 and K_2O contents in the sanidine-bearing system and the polyphase composition of the subsolidus (increasing number of phases decreases the melting temperature for the eutectic). Relatively low temperature of the Cpx and Grt equilibria with the K–Al–Si-rich melts is in good agreement with mineral thermometry of deep-seated assemblages, which in the most cases indicate temperatures 1,250–1,300°C. Addition of Na, Fe, and other components would further lower temperatures for coexistence of Cpx , Grt , and Opx with K-rich melts at pressures of about 6 GPa.

Some natural eclogite assemblages bear evidence for equilibrium with potassium-rich aluminosilicate melts (Bulanova and Argunov 1985; Bulanova et al. 1988; Novgorodov et al. 1990), whose compositions were close to the melts that we have described in the system $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 . One of the most important

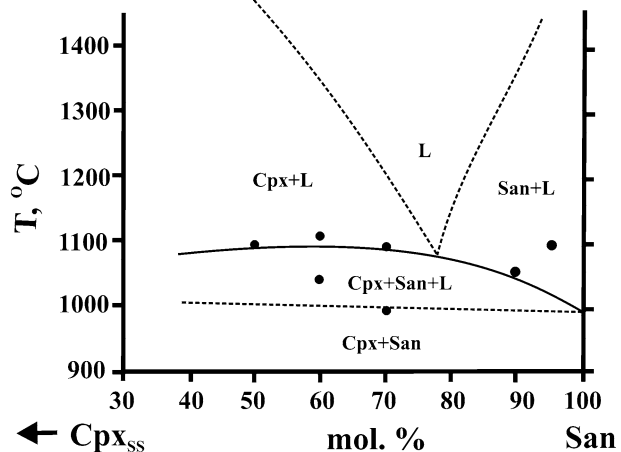


Fig. 7 Preliminary T–X phase diagram for the system $\text{CaMgSi}_2\text{O}_6$ – KAlSi_3O_8 (join Cpx_{ss} – San) at 3.5 GPa. Solid lines denote experimentally established boundaries of phase fields; dashed lines indicate the assumed boundaries of phase fields

features of the natural assemblages is the concentration of K_2O in clinopyroxenes, which varies from 0.2 wt% in some eclogite xenoliths up to 1.4 wt% in clinopyroxene inclusions in diamonds (e.g. Stachel et al. 2000a). Using the negative effect of jadeite on the KJd solubility in the UHP Cpx, we demonstrated previously (Safonov et al. 2004) that the potassium-bearing omphacite inclusions in diamonds could crystallize from the potassium-rich melts at pressures below 6.5 GPa. Despite the presently low K_2O content, many Cpx from eclogite xenoliths also show evidence for initially elevated potassium content, such as sanidine lamellae (Reid et al. 1976; Schulze and Helmstaedt 1988; Smith et al. 1991; Schmickler et al. 2004). Some melt inclusions in diamonds from kimberlites (Prinz et al. 1975; Bulanova and Argunov 1985; Bulanova et al. 1988; Novgorodov et al. 1990; Navon et al. 2003; Shiryayev et al. 2003) show high contents of SiO_2 (61–67 wt%), Al_2O_3 (7–21 wt%), and $K_2O + Na_2O$ (6–18 wt%) at low concentration of CaO, MgO, FeO, and volatile components. These concentrations are also characteristics of the melts along the $CaMgSi_2O_6$ – $KAlSi_3O_8$ join (Table 2). Inclusions of potassic aluminosilicate melts are commonly accompanied by Cpx, Grt, and San (Bulanova and Argunov 1985; Bulanova et al. 1988; Novgorodov et al. 1990), which are also major phases in the join $CaMgSi_2O_6$ – $KAlSi_3O_8$ (Tables 1 and 2). These data allow application of our experimental results to reconstruction of the evolution of these natural assemblages.

Here we consider theoretically just two specific features of the evolution of potassic melts and their assemblages deduced from the present experimental results: (1) the relationships between KCpx and San and (2) the sequence of mineral crystallization from the K-rich melts. The detailed application of the experimental results for the K-rich aluminosilicate systems to natural assemblages will be discussed in subsequent publications.

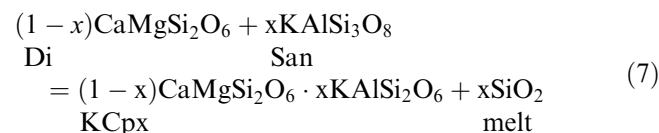
Relations between potassium-bearing clinopyroxene and sanidine

The phase diagram in Fig. 6 shows that only melts with $N_{San}^{system} > 70$ mol% are able to produce the assemblage Cpx + San at pressures close to the upper boundary of the sanidine stability. At the lower N_{San}^{system} , the Cpx + San assemblage is unstable because of the complex reactions (4)–(6), and Cpx is the only solid phase hosting potassium (Figs. 4 and 6).

In the San-bearing portion of the system, KCpx appears at $\sim 1,300^\circ C$ (experiment 1046; Fig. 6). The value $N_{KJd}^{Cpx} \sim 6$ mol% roughly agrees with N_{KJd}^{Cpx} of San-free experiments at similar temperature (Fig. 4). The formation of almost pure Di associated with San in the experiment 1086 at similar temperature (Table 2) is, probably, related to shielding of diopside by large San crystals (Fig. 1e). This interpretation is readily

supported by the presence of KCpx and diopside in this experiment, as well as in both the 1087 and 1194. KCpx forms separate crystals or clusters among the quench products. Thus, KCpx is in equilibrium with the melt. Similar texture, produced by separation of phases in the melt (see above), is characteristic also for the garnet-bearing experiments containing both KCpx and diopside (experiments 1089, 1090, 1121).

The experiments 1046, 1086, 1087, and 1194 display textures that are consistent with reaction relationships between Cpx, San, and melt (Figs. 1e and 2d). Figure 2d shows a partial replacement of San by the aggregate of KCpx (N_{KJd}^{Cpx} up to 4.9 mol%) and quenched melt. Relics of semi-reacted diopside are locally preserved inside the KCpx + Q aggregates (Fig. 2d). The above texture corresponds to the following reaction:



However, no reaction is observed between shielded Di and San (Fig. 2c). This observation allows the conclusion that the formation of KCpx after San is not a solid-state reaction: it proceeds only in the presence of melt. Therefore, the reaction (7) can be rewritten with $CaMgSi_2O_6$ as a component of the coexisting melt. The large (up to 300 μm) San crystals and cumulates contain “pockets” of the quenched melt consisting of acicular sanidine and residual glass (Figs. 1f and 2c). The composition of this quench, however, is very different from the composition of the melt outside the cumulates (see Table 2 for analyses of glass and Q in the experiment 1090). The major difference is the higher SiO_2 content in the quenched melt from the “pockets” (70 vs. 56 wt%). Thus, at high SiO_2 content in the melt (high silica activity) at constant P and T , the reaction (7) is strongly displaced to the left (see below).

Thus, our results are consistent with the suggestion that the formation of KCpx in the sanidine-bearing portion of the system $CaMgSi_2O_6$ – $KAlSi_3O_8$ (Fig. 6) proceeded via the crystals-melt reaction (7). At $x=1$ this reaction could be re-written as $San = KJd + L$, which has a form of a peritectic reaction. Although KCpx is formed as a result of the reaction (7), San does not totally react out. The Cpx composition is controlled by the solidus line in the $CaMgSi_2O_6$ -rich portion of the system, which is roughly drawn in accordance with the general dependence of the Cpx composition on temperature (Fig. 4) and defines a maximal solubility of KJd in Cpx on isotherms at 6 GPa. Figure 6 indicates that in the sanidine-bearing portion of the join first KCpx appears via the reaction (7), if the sanidine liquidus crosses the isotherm about $1,300^\circ C$. The isotherm corresponds to ~ 5 mol% of KJd in Cpx. Therefore, this value is a minimal N_{KJd}^{Cpx} in the equilibrium $San + KCpx + L$ at 6 GPa. Below $1,300^\circ C$, KCpx is stable with San and melt, and N_{KJd}^{Cpx} increases systematically with decreasing

relatively low temperatures (Fig. 6), and its amount in a rock would be significantly subordinate to Cpx. No San appears at HP in the melts at $N_{\text{San}}^{\text{system}} < 70$ mol% (Fig. 6). However, at $P < 4$ GPa, garnet does not crystallize, whereas San forms from the CaMgSi₂O₆-rich melts. The residual melt would crystallize to the eutectic assemblage of K-poor Cpx and San (Fig. 7, Perchuk and Yapaskurt 1998).

This scheme of crystallization was proposed by Perchuk et al. (2002, 2003) and Bindi et al. (2003) for the Ca-Fe rock varieties of Grt-Cpx rocks of the Kokchetav Complex, N. Kazakhstan, where Mg-number of clinopyroxene is below 55 mol%. However, in order to closely apply the results of the experiments in the Di-San join, we have to consider another rock variety, i.e. Mg-rich Grt-Cpx rocks of the same complex, where Mg-number of Cpx reaches 90 mol% (e.g. Sobolev and Shatsky 1990). Grt is quantitatively and texturally subordinate to Cpx, forming small rounded grains at the boundaries and inclusions in the peripheral zones of large Cpx crystals. In turn, Cpx inclusions in Grt are very rare. These inclusions contain variable concentration of K₂O (0.4–0.7 wt%). The K₂O content in the KCpx decreases from cores to rims of inclusions. Rare separate inclusions as well as cores of large Cpx crystals in matrix are crowded with K-feldspar lamellae. Later K- and lamellae-free Cpx form rims around the matrix Cpx. In addition to lamellae in Cpx, K-feldspar forms separate intergranular clusters between clinopyroxene and garnet grains.

Thus, the textural and compositional characteristics of the Mg-rich garnet-clinopyroxene rocks are very similar to the crystallization sequence for the melts in the diopside-rich portion of the join CaMgSi₂O₆-KAlSi₃O₈. Since primary San is absent in these rocks, we assume that crystallization of the rocks began at $P > 4$ GPa. KCpx with variable K₂O contents crystallized during cooling and decompression. Grt joined Cpx subsequently. Because Grt was crystallized at a relatively later stage among larger clinopyroxene grains, the possibility of Cpx entrapment by Grt is very low. The wide variations of K₂O content in KCpx inclusions, probably, points to their entrapment at very rapid cooling (resulting in increase of the potassium content) and decompression (resulting in decrease of the potassium content). Coexistence of potassium feldspar with K-poor Cpx corresponds to a final stage of magmatic evolution at $P < 4$ GPa.

The Cpx contents in both the Cpx-Grt-San and Cpx-Opx-Grt-San xenoliths from kimberlites are also above 40 vol% (Reid et al. 1976; Schmickler et al. 2004). Small grains of Grt, Opx, and San occur at the boundaries of large clinopyroxene crystals, and lamellae in Cpx crystals (e.g. Schmickler et al. 2004). All these textural features correspond to crystallization of Grt, Opx, and San later than Cpx and agree with the above scheme of crystallization from aluminosilicate melts of the diopside-rich portion of the join CaMgSi₂O₆-KAlSi₃O₈.

The characteristic feature of the above rocks is cumulate-like texture, which implies a fractionation of

minerals during the crystallization process. Cumulates are composed of clinopyroxene and garnet, whereas sanidine and other K-Si-rich minerals occupy intergranular aggregates. Textures in some experiments of the join CaMgSi₂O₆-KAlSi₃O₈ (Fig. 2a–c) suggests a possibility of gravitational phase separation in the K-rich aluminosilicate melts, which produces Cpx + Grt cumulates cemented with the quenching products of the melt. We previously noted similar Cpx + Grt accumulation in the experiments for the join CaMgSi₂O₆-KAlSi₂O₆ at 7 GPa (Safonov et al. 2002, 2003). The major characteristics of melts both in the present join and in the join CaMgSi₂O₆-KAlSi₂O₆ are high concentration of SiO₂, Al₂O₃ and K₂O. It is well known that the viscosity of highly polymerized at ambient pressure melts rapidly *decreases* by orders of magnitude with pressure (e.g., Kushiro 1981; Persikov 1991). Therefore, the gravitational mechanism of Cpx, Grt, and San separation is physically substantiated for the UHP conditions. Our experimental results visualize this mechanism.

According to our results (Figs. 6 and 7), early (liquidus) crystallization of San at 6 GPa is possible only from very K-Al-Si rich melts. The rare findings of primary inclusions of San or its intergrowths with Cpx or Grt in diamonds (Prinz et al. 1975; Meyer and McCallum 1986; Bulanova and Argunov 1985; Novgorodov et al. 1990; Sobolev et al. 1998) correlates with scarcity of the K-rich aluminosilicate melt inclusions in diamonds (Prinz et al. 1975; Bulanova and Argunov 1985; Bulanova et al. 1988; Novgorodov et al. 1990; Shiryayev et al. 2003; Navon et al. 2003). Probably, the formation of San-producing melts at HP conditions is related to the late-stage evolution of more abundant mantle potassium-rich liquids of the carbonate-aluminosilicate-brine system (e.g., Navon et al. 2003 and references therein).

Conclusions

This experimental study of the join CaMgSi₂O₆-KAlSi₃O₈ at 6 and 3.5 GPa allows the following major conclusions.

1. Melts with compositions close to the system CaMgSi₂O₆-KAlSi₃O₈, can produce a wide spectrum of eclogite-like assemblages involving Cpx, Grt, Opx, and San. Because of high K₂O and SiO₂ contents, these melts coexist with Grt, potassium-bearing Cpx, and San down to ~1,100°C. This conclusion allows substantiation of a magmatic origin for mantle parageneses containing potassium-bearing Cpx or Cpx + San at $T < 1,200^\circ\text{C}$.
2. Potassium-bearing Cpx or San are two major competing K-bearing phases within the “dry” UHP assemblages. San is unstable in the eclogitic (diopside-rich) melts at $P \sim 6$ GPa, being replaced by the assemblage of potassium-bearing Cpx (\pm Opx) and

Grt. In addition to the San instability in the peridotite systems (e.g., Wendlandt and Egger 1980), this could be a major reason for extreme scarcity of San in the mantle assemblages. In contrast, Cpx becomes a major host of K_2O at P down to about 4 GPa.

- Liquidus crystallization of San and its assemblage with the KCpx and Grt at the UHP environments corresponds to the evolution of the extremely potassic and silica-rich aluminosilicate melts. These conclusions characterize the ultrapotassic melts as active participants of magmatic processes in the mantle, which dramatically change thermodynamic conditions of the formation of the “ordinary” deep-seated assemblages.

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